

INTRODUCTION

Bedrock of the Long Branch quadrangle consists of unconsolidated sand, silt, clay, and glauconite clay laid down in coastal, nearshore-marine, and continental-shelf settings between 95 and 10 million years ago. The sediments are classed to 19 formations and members. Lithology and age of the formations are provided in the *Description of Map Units*. Age of the formations and their bounding unconformities are summarized in the *Correlation of Map Units*. Cross sections AA', BB', and CC' show the subsurface geometry of the formations along the line of section. Surficial sand and gravel, silt sand, and organic silt and clay of the Miocene, Pliocene, and Quaternary age, overlie the bedrock in most of the quadrangle. The surficial deposits include fluvial, estuarine, and nearshore-marine sediments and were mapped by Stanford (2000). They are shown by overprint pattern on the map where they are more than 5 feet thick.

DESCRIPTION OF MAP UNITS

- Tch** COHANSEY FORMATION—Fine-to-course quartz sand, with thin beds of very coarse sand to very fine pebbles; very pale brown, white, light gray. Weakly horizontally bedded to cross-bedded. Sand and very fine pebbles consist of quartz with minor (<5%) quartzite and chert. Course sand beds are locally iron-cemented. As much as 30 feet thick in the Long Branch quadrangle. In hilltop cross-section remnants above 140-160 feet in elevation in the southwest corner of the quadrangle. Latest Middle Miocene in age, based on pollen (Owens and others, 1988, 1998). Unconformably overlies the Kirkwood Formation.
- Thw** KIRKWOOD FORMATION—Very-fine-to-fine quartz sand, minor medium-to-course sand, with thin interbeds of clay and silt; micaceous and lignitic, with a trace (<1%) of glauconite; white and light gray where unweathered, dark gray where unweathered. Sand is unstratified to horizontally bedded. In subsurface only. As much as 240 feet thick in the eastern part of the quadrangle, thins to 160 feet thick in the central and western parts of the quadrangle. Late Cretaceous (early to middle Campanian) in age based on pollen (Wolfe, 1976), ostracodes (Gohn, 1992), calcareous nanofossils, and strontium-isotope ratios (Miller and others, 2006). Grades downward into the Merchantville Formation. On geophysical well logs, transition to Woodbury is marked by increased gamma-ray intensity and decreased resistance.
- Tmq** MANASQUAN FORMATION—Glauconitic (15-30%) clayey-silty very-fine-to-fine sand to fine-sandy clayey silt; olive and olive-brown where unweathered, yellowish-brown where weathered; unstratified. As much as 80 feet thick. Early Eocene in age based on calcareous nanofossils (Sugarmann and others, 1995; Owens and others, 1998; Miller and others, 2006) and foraminifera (Miller and others, 2006). Described by drillers as "green sand" or "green clay". Unconformably overlies the Vincentown Formation.
- Tvt** VINCENTOWN FORMATION—Glauconitic (5-20%) silty medium-to-course quartz sand, some fine-to-medium sand, some very coarse sand to very fine pebbles; yellow, reddish-yellow, olive-brown, unstratified to weakly horizontally stratified. Coarse sands are locally iron-cemented into beds and masses as much as 10 feet thick. Lowermost 10-20 feet of the formation is silty fine-to-medium sand, with more glauconite than upper part. Total thickness of formation is 180 feet. Late Paleocene in age, based on foraminifera (Olsson and Wise, 1987; Miller and others, 2006) and calcareous nanofossils (Sugarmann and others, 1991). Unconformably overlies the Hornerstown Formation.
- Th** HORNERSTOWN FORMATION—Glauconitic (>50%) clay and silty clay; olive, dark brown, black where unweathered, olive-brown to reddish-brown mottles where weathered; unstratified. Glauconite occurs primarily in soft grains of fine-to-medium sand size. Thickness is 25 to 30 feet. Early Paleocene in age, based on foraminifera (Olsson and others, 1997; Landman and others, 2004; Miller and others, 2006) and calcareous nanofossils (Sugarmann and others, 1991; Miller and others, 2006). Unconformably overlies the Tinton Formation.
- Kt** TINTON FORMATION—Glauconitic (5-30%) silty medium-to-course and fine-to-medium quartz sand, reddish-brown, reddish-yellow, yellowish-brown where weathered, grayish-brown, brown, olive-brown where unweathered; unstratified to weakly horizontally stratified. Commonly iron-cemented into beds and masses as much as 15 feet thick. Uppermost 4-6 feet, just below contact with Hornerstown Formation, is a transition to olive-gray glauconitic clayey-silty fine to fine-sandy silt ("New Egypt Formation" of Landman and others, 2004). Total thickness of Tinton is 30 to 40 feet. Late Cretaceous (late Maestrichtian) in age based on foraminifera, nanofossils, and ammonites (Landman and others, 2004) and strontium-isotope ratios (Sugarmann and others, 1995). Overlies the Shrewsbury Member of the Red Bank Formation. Contact with Shrewsbury is not exposed in the Long Branch quadrangle. It is gradational over several feet in the Sandy Hook quadrangle, north of the Long Branch quadrangle (Minard, 1969), but may be unconformable in the Marlboro quadrangle, west of the Long Branch quadrangle (Sugarmann and Owens, 1998).
- Krs** RED BANK FORMATION, SHREWSBURY MEMBER—Fine-to-medium quartz sand, minor medium-to-course sand, slightly silty, glauconitic (<5%), and micaceous; reddish-yellow, yellow where weathered, light gray and gray where weathered; unstratified to weakly horizontally bedded, locally iron-cemented. As much as 100 feet thick. Late Cretaceous (late Maestrichtian) in age based on fossils in the underlying Sandy Hook Member, the Shrewsbury Member is unfossiliferous. Grades downward within 2-3 feet to the Red Bank Formation, Sandy Hook Member. On geophysical well logs, transition to Sandy Hook Member is marked by increased gamma-ray intensity and decreased resistance.
- Krs** RED BANK FORMATION, SANDY HOOK MEMBER—Fine-sandy clayey silt, micaceous, slightly glauconitic (<5%), brown to yellowish-brown where unweathered, dark gray, olive-gray where unweathered; unstratified. Calcareous brachiopod, pelecypod, and gastropod fossils are common. As much as 20 feet thick. Late Cretaceous (late Maestrichtian) in age based on calcareous nanofossils (Sugarmann and Owens, 1996), foraminifera (Olsson, 1964; Olsson and Wise, 1987; Owens and others, 1977), and strontium-isotope ratios (Sugarmann and others, 1995). Grades downward within 2-3 feet into the Navasink Formation. On geophysical well logs, transition to Navasink is marked by increased gamma-ray intensity and slightly decreased resistance.
- Kns** NAVASINK FORMATION—Glauconitic (30-50%) clayey-silty fine-to-medium quartz sand to fine-sandy clayey silt; dark gray, gray where unweathered, brown to yellowish-brown where weathered. Glauconite occurs chiefly in soft grains of fine-to-medium sand size. Calcareous brachiopod, pelecypod, and gastropod fossils are common. Late Cretaceous (late Maestrichtian) in age based on calcareous nanofossils and foraminifera (Olsson, 1964; Miller and others, 2006), macrofossils (Sohl, 1977), and strontium-isotope ratios (Sugarmann and others, 1995). Unconformably overlies the Mount Laurel Formation. Contact with Mount Laurel is commonly marked by a sharp peak in gamma-ray intensity on geophysical well logs, with reduced intensity in the Mount Laurel.
- Kml** MOUNT LAUREL FORMATION—Glauconitic (3-15%) fine-to-medium quartz sand, minor medium-to-course sand, with thin interbeds of clay and silt; yellowish-brown where weathered, olive-gray to olive-brown where unweathered. Sand is unstratified to horizontally bedded to cross-bedded. As much as 50 feet thick in the southern part of the quadrangle; thins to 20 feet to the north. In subsurface only, covered by surficial deposits in the Navasink River estuary and by overlying Coastal Plain formations elsewhere. Late Cretaceous (late Campanian) in age, based on calcareous nanofossils and strontium-isotope ratios (Sugarmann and others, 1991; Miller and others, 2006). Grades downward into the Wenonah Formation. On geophysical well logs, transition to Wenonah is generally marked by slightly decreased resistance and increased gamma-ray intensity.
- Kw** WENONAH FORMATION—Silty fine-to-very-fine quartz sand to fine-sandy clayey silt, micaceous, slightly glauconitic (<5%), yellow, very pale brown where unweathered, gray to pale-olive where unweathered, unstratified. As much as 40 feet thick. In subsurface only, covered by surficial deposits in the Navasink River estuary and by overlying Coastal Plain formations elsewhere. Late Cretaceous (late Campanian) in age based on pollen (Wolfe, 1976) and ammonites (Kennedy and Cobban, 1994). Grades downward into the Marshalltown Formation. On geophysical well logs, transition to Marshalltown is marked by increased gamma-ray intensity.

- Km** MARSHALLTOWN FORMATION—Glauconitic (20-50%), slightly micaceous, silty, clayey fine-to-medium quartz sand, to fine-sandy clayey silt; olive-gray to olive-brown; unstratified. Thickness is 14 to 20 feet. In subsurface only. Late Cretaceous (middle Campanian) in age based on calcareous nanofossils, foraminifera, mollusks, and strontium-isotope ratios (Sugarmann and others, 1995). Unconformably overlies the Englishtown Formation. On geophysical well logs, contact with Englishtown is marked by decreased gamma-ray intensity and slightly increased resistance.
 - Ke** ENGLISHTOWN FORMATION—Fine-to-medium quartz sand, minor medium-to-course sand, with thin interbeds of clay and silt; micaceous and lignitic, with a trace (<1%) of glauconite; white and light gray where unweathered, dark gray where unweathered. Sand is unstratified to horizontally bedded to cross-bedded. In subsurface only. As much as 140 feet thick in the eastern part of the quadrangle, thins to 110 feet thick in the west. In the Asbury Park quadrangle to the south of the Long Branch quadrangle, and farther southwest in northern Ocean County, the Englishtown is divided into an upper and lower member based on the presence of a clay-silt facies in the middle of the formation that is distinctive on gamma-ray logs (Nichols, 1977; Sugarmann and Owens, 1994; Miller and others, 2006). This facies is not well marked on gamma-ray logs in the Long Branch quadrangle (wells 29-2935, 29-2941, 29-2946, 29-4673, 29-2348, and 29-48307) and so the members are not mapped here. Late Cretaceous (middle to late Campanian) in age, based on pollen (Wolfe, 1976), ostracodes (Gohn, 1992), calcareous nanofossils, and strontium-isotope ratios (Miller and others, 2006). Grades downward into the Woodbury Formation. On geophysical well logs, transition to Woodbury is marked by increased gamma-ray intensity and decreased resistance.
 - Kwb** WOODBURY FORMATION—Clay, silty clay, with minor thin beds of very fine quartz sand, slightly micaceous and lignitic; dark gray and black where unweathered, yellowish-brown to brown where weathered; unstratified. In subsurface only. As much as 240 feet thick in the eastern part of the quadrangle, thins to 160 feet thick in the central and western parts of the quadrangle. Late Cretaceous (early to middle Campanian) based on pollen (Wolfe, 1976), ostracodes (Gohn, 1992), and calcareous nanofossils (Miller and others, 2006). Grades downward into the Merchantville Formation. On geophysical well logs, transition to the Merchantville is marked by slightly increased gamma-ray intensity.
 - Kmg** MERCHANTVILLE FORMATION—Glauconitic (20-50%) clayey silt to sandy clayey silt, slightly micaceous, olive, dark gray, black where unweathered, olive-brown to yellowish-brown where weathered; unstratified. Thickness is 40 to 60 feet. In subsurface only. Late Cretaceous (early Campanian to Santonian) in age based on ammonites (Owens and others, 1977) and calcareous nanofossils (Miller and others, 2006). The Chesapeake Formation, a glauconitic clayey silt underlying the Merchantville, is mapped in northern Monmouth and eastern Middlesex counties (Sugarmann and Owens, 1996; Sugarmann and others, 2005; Stanford and Sugarmann, 2006) and in the subsurface both west and south of the Long Branch quadrangle (Sugarmann and Owens, 1994, 1996). Because it is lithically similar to the Merchantville and cannot be easily distinguished from it on geophysical logs, it is not mapped separately. If present, it is included here within the Merchantville, or uppermost Magogy Formation.
 - Km** MAGOXY FORMATION—Fine-to-medium quartz sand, some very-fine-to-fine sand and minor medium-to-course sand, micaceous, lignitic, and pyrite-bearing in places, with thin interbeds of silt and clay, white to yellow where unweathered, light gray to gray where unweathered. Sand is cross-bedded to laminated. As much as 220 feet thick. In subsurface only. Late Cretaceous (Turonian-Santonian) in age, based on pollen (Christopher, 1979, 1982; Miller and others, 2006). Unconformably overlies the Raritan Formation, Woodbridge Clay member. On geophysical well logs, contact with the Woodbridge is marked by increased gamma-ray intensity.
- In its outcrop area in eastern Middlesex County the Magoxy is divided into 5 members. From bottom to top they include: South Asbury Fire Clay, Old Bridge Sand, Asbury Stoneare Clay, Morgan beds, and Cliffwood beds (Sugarmann and others, 2005). The Old Bridge is a thick sand, the other members are interbedded clay-silt and fine sand. These members may extend downward in the subsurface (Miller and others, 2006). Geophysical well logs in the Long Branch quadrangle (wells 29-2162, 29-2348, 29-2151, 29-2935, 29-2941, 29-2946, and 29-4173) show generally higher gamma-ray intensity and lower resistivity in the uppermost 50 feet of the formation, and again in the lowermost 30-40 feet, than in the middle 100-120 feet. The upper fine-grained beds may correspond to the Asbury Stoneare Clay and Morgan and Cliffwood beds, and the lower fine-grained beds may correspond to the South Asbury Fire Clay. The middle sand may correspond to the Old Bridge.
- Kw** RARITAN FORMATION, WOODBRIDGE CLAY MEMBER—Clay and silt, micaceous, lignitic, and pyrite-bearing; gray and black where unweathered, white to brown where weathered; with minor thin interbeds and laminae of white, yellow, and light gray very-fine-to-fine quartz sand. As much as 110 feet thick. In subsurface only, penetrated by wells 29-2465 and 29-1921. The driller's log for well 29-2465 in Easton reports "weathered bedrock", with no further information, beneath the Magoxy Formation, at a depth of 875-891 feet. This depth is anomalously shallow for the basement surface, suggesting that the material may be weathered clay of either the Woodbridge or South Asbury Fire Clay member of the Magoxy. The Woodbridge is Late Cretaceous (late Cenomanian) in age based on pollen (Christopher, 1979) and ammonites (Cobban and Kennedy, 1990). Grades downward into the Raritan Formation, Farrington Sand member. Transition to the Farrington is marked by decreased gamma-ray intensity on geophysical well logs.
 - Krf** RARITAN FORMATION, FARRINGTON SAND MEMBER—Fine-to-course quartz sand, some coarse-to-very-course sand, minor beds of clay and silt; white and yellow where unweathered, gray where unweathered. Sands are horizontally bedded to cross-bedded. As much as 60 feet thick. In subsurface only, penetrated in well 29-2465. Late Cretaceous (Cenomanian) in age based on pollen (Christopher, 1979). Unconformably overlies the Potomac Formation. Contact with Potomac is marked by increased gamma-ray intensity on geophysical well logs.
 - Kp** POTOMAC FORMATION—Fine-to-medium quartz sand, some coarse-to-very-course sand, with beds of clay and silt; white and yellow where unweathered, gray where unweathered. Sands are horizontally bedded to cross-bedded, clay is in the Long Branch quadrangle. In subsurface only, partially penetrated in well 29-2465. Late Cretaceous (Albian-Cenomanian) based on pollen (Sugarmann and Owens, 1996; Miller and others, 2006), which indicates that the Potomac in this area corresponds to the Potomac Formation, unit 1, of Doyle and Robbins (1977).

MAP SYMBOLS

- T** Contact—Approximately located. Solid triangle indicates contact observed in outcrop. Open triangle indicates contact formerly observed, as reported in permanent note collection of the N. J. Geological Survey.
 - o** Formation observed in outcrop or excavation, or penetrated in hand-auger hole.
 - o** Formation formerly observed in outcrop or excavation—Reported in permanent note collection of the N. J. Geological Survey.
 - o** Formation covered by surficial deposits—Surficial deposits of Quaternary, Pliocene, and late Miocene age continuous and generally more than 5 feet thick.
- Well showing formations penetrated—Location accurate to within 200 feet. Identifiers of the form 29-xxxx are N. J. Department of Environmental Protection well permit numbers. Identifiers of the form 25-xxxx are U. S. Geological Survey Ground Water Site Inventory identification numbers. Lithologic and geophysical logs for most of these wells are provided by Goussier and others (1989). Identifiers of the form 29-xxxx are N. J. Atlas Sheet coordinates of records of wells in the permanent note collection of the N. J. Geological Survey. Identifiers of the form "29-2343 B1" provide data and identification number of test borings drilled by the A. J. Healy Company, with copies on file at the N. J. Geological Survey. Identifier "KPI" indicates a stratigraphic test hole drilled by the U. S. Geological Survey in cooperation with the N. J. Geological Survey and Rutgers University Earth and Planetary Sciences Department in 2008.

Number preceding formation symbol is depth, in feet below land surface, of base of unit, as inferred from lithologic logs. Final number is total depth of well or boring rather than base of unit. "Surficial" indicates surficial deposits. Units joined with a "-" cannot be separately identified in the drillers' descriptions. For wells shown on sections the formations penetrated are not listed on the map. Drillers' logs vary in detail and accuracy. They are used in combination with outcrop data and geophysical well logs to map contacts, so depths of some contacts inferred from the logs may not match those shown on the map and sections.

Well showing formations penetrated—Location accurate to within 500 feet. Identifiers and log information as above.

Geophysical log—On sections, Gamma-ray log is indicated by "G" after well identifier, and is shown by a red line. Intensity of gamma-ray radiation increases to right. Electric log is indicated by "E" after well identifier and is shown by paired blue lines, with spontaneous potential shown on left-hand curve (voltage increasing to right) and resistivity shown on right-hand curve (resistance increasing to right). Horizontal scale varies from well to well. Arrows indicate gamma responses that are due to well construction features rather than formation lithology (well 29-2935, section AA' only).

REFERENCES

Christopher, R. A., 1979, Normapollens and triporate pollen assemblages from the Raritan and Magoxy Formations (Upper Cretaceous) of New Jersey. *Palynology*, v. 3, p. 73-121.

Christopher, R. A., 1982, The occurrence of the *Complanatopsis-titanopolis* Zone (polynormans) in the Eagle Ford Group (Upper Cretaceous) of Texas. *Journal of Paleontology*, v. 56, p. 525-541.

Cobban, W. A., and Kennedy, W. J., 1990, Upper Cenomanian ammonites from the Woodbridge Clay member of the Raritan Formation in New Jersey. *Journal of Paleontology*, v. 64, p. 845-846.

Doyle, J. A., and Robbins, E. I., 1977, Angiosperm pollen zonations of the Cretaceous of the Atlantic Coastal Plain and its application to deep wells in the Salisbury embayment. *Palynology*, v. 1, p. 47-78.

Gohn, G. S., 1992, Preliminary ostracode biostratigraphy of subsurface Campanian and Maestrichtian sections of the New Jersey Coastal Plain. In Gohn, G. S., ed., *Proceedings, U. S. Geological Survey workshop on the geology and geology of the Atlantic Coastal Plain, 1988*. U. S. Geological Survey Circular 1059, p. 15-21.

Gronberg, J. M., Birkelo, B. A., and Pucci, A. A., 1989, Selected benthic geology of the Atlantic Coastal Plain and its application to deep wells in the Salisbury embayment. *Palynology*, v. 1, p. 47-78.

Kennedy, W. J., and Cobban, W. A., 1994, Ammonite faunas from the Woonah Formation (Upper Cretaceous) of New Jersey. *Journal of Paleontology*, v. 68, p. 9-11.

Landman, N. H., Johnson, R. O., and Edwards, L. E., 2004, Cephalopods from the Cretaceous-Tertiary boundary interval on the Atlantic Coastal Plain, with a description of the highest ammonite zones in North America, part 2, northeastern Monmouth County, New Jersey. *Bulletin of the American Museum of Natural History*, number 287, 107 p.

Miller, K. G., Sugarmann, P. J., Browning, J. V., Aubrey, M. P., Bremer, G. J., Cobbs, G., III, deRomero, L., Feigenson, M. D., Harris, A., Katz, M. E., Kalpoe, A., McLaughlin, P. J., Minetteva, S., Monahan, D. H., Olsson, R. K., Patrick, L., Pekar, S. J., and Upretige, J., 2006, Sea grid site, in Miller, K. G., Sugarmann, P. J., and Browning, J. V., eds., *Proceedings of the Ocean Drilling Program, Initial Reports*, v. 174A, p. 1-104.

Minard, J. P., 1969, Geology of the Sandy Hook quadrangle in Monmouth County, New Jersey. U. S. Geological Survey Bulletin 1276, 43 p.

Nichols, W. D., 1977, Digital computer simulation model of the Englishtown aquifer in the northern Coastal Plain of New Jersey. U. S. Geological Survey Water Resources Investigations WRI 77-73, 101 p.

Olsson, R. K., 1964, Late Cretaceous planktonic foraminifera from New Jersey and Delaware. *Journal of Paleontology*, v. 10, p. 157-188.

Olsson, R. K., Miller, K. G., Browning, J. V., Habib, D., and Sugarmann, P. J., 1997, Ejecta layer at the K/T boundary, Bass River, New Jersey (ODP Leg 174AX). *Geology*, v. 25, p. 759-762.

Olsson, R. K., and Wise, S. W., Jr., 1987, Upper Paleocene to middle Eocene depositional sequences and hiatuses in the New Jersey Atlantic margin. In Ross, C. A., and Haman, D., eds., *Tuning and depositional history of eustatic sequences—constraints in seismic stratigraphy*. Publication Foundation for Foraminiferal Research Special Publication 24, p. 99-112.

Owens, J. P., Bykirk, L. M., Paulschok, G., Agat, T. A., Gonzalez, V. M., and Sugarmann, P. J., 1988, Stratigraphy of the Tertiary sediments in a 945-foot corehole near Mays Landing in the southeastern New Jersey Coastal Plain. U. S. Geological Survey Professional Paper 1484, 39 p.

Owens, J. P., Sobel, N. F., and Minard, J. P., 1971, A field guide to Cretaceous and lower Tertiary beds of the Raritan and Salisbury embayments, New Jersey, Delaware, and Maryland. *American Association of Petroleum Geologists Society of Economic Paleontologists and Mineralogists*, 113 p.

Owens, J. P., Sugarmann, P. J., Sobel, N. F., Parker, R. A., Hoaghton, H. F., Volkert, R. A., Duke, A. A., Jr., and Omdorf, C. C., 1998, Bedrock geologic map of central and northern New Jersey. *Geological Survey Miscellaneous Investigations Series Map 1-2540-B*, scale 1:100,000.

Sobel, N. F., 1977, Benthic marine molluscan associations from the Upper Cretaceous of New Jersey and Delaware. In Owens, J. P., Sobel, N. F., and Minard, J. P., eds., *A field guide to Cretaceous and lower Tertiary beds of the Raritan and Salisbury embayments, New Jersey, Delaware, and Maryland*. American Association of Petroleum Geologists-Society of Economic Paleontologists and Mineralogists, p. 70-91.

Stanford, S. D., 2000, Surficial geology of the Long Branch quadrangle, Monmouth County, New Jersey. N. J. Geological Survey Open-File Map OFM 38, scale 1:24,000.

Stanford, S. D., and Sugarmann, P. J., 2008, Bedrock geology of the Jamesburg quadrangle, Middlesex, Monmouth and Mercer counties, New Jersey. N. J. Geological Survey Open-File Map OFM 72, scale 1:24,000.

Sugarmann, P. J., Miller, K. G., Bakry, D., and Feigenson, M. D., 1995, Uppermost Campanian-Maestrichtian strontium isotope, biostratigraphic, and sequence stratigraphic framework of the New Jersey Coastal Plain. *Geologic Society of America Bulletin*, v. 107, p. 19-37.

Sugarmann, P. J., Miller, K. G., Owens, J. P., and Feigenson, M. D., 1993, Strontium isotope and sequence stratigraphy of the Miocene Kirkwood Formation, southern New Jersey. *Geological Society of America Bulletin*, v. 105, p. 423-436.

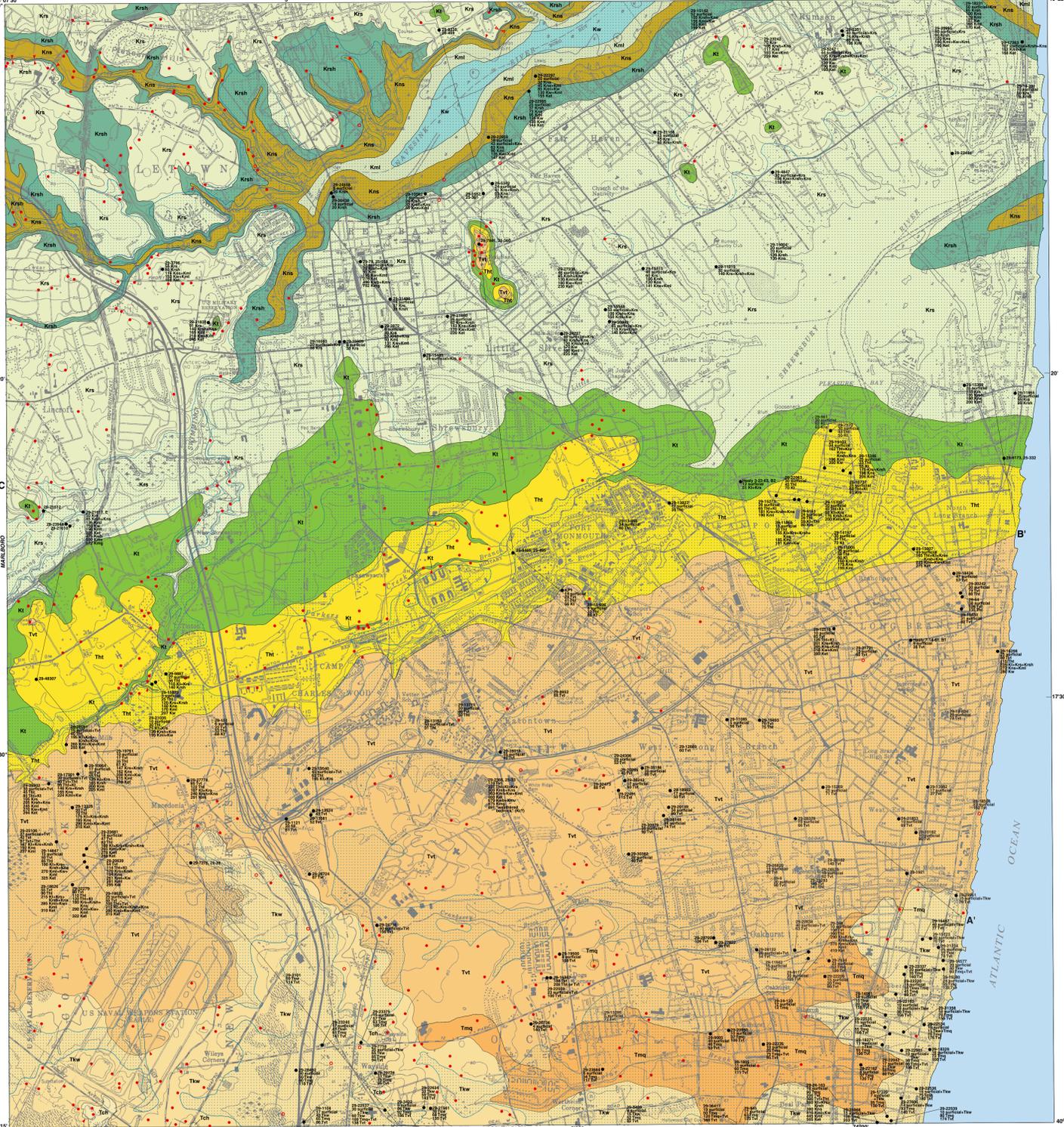
Sugarmann, P. J., and Owens, J. P., 1994, Geologic map of the Asbury Park quadrangle, Monmouth and Ocean counties, New Jersey. N. J. Geological Survey Geologic Map Series GMS 94-2, scale 1:24,000.

Sugarmann, P. J., and Owens, J. P., 1996, Bedrock geologic map of the Freehold and Marlboro quadrangles, Middlesex and Monmouth counties, New Jersey. N. J. Geological Survey Geologic Map Series GMS 96-1, scale 1:24,000.

Sugarmann, P. J., Owens, J. P., and Bybell, L. M., 1991, Geologic map of the Adelphi and Farrington quadrangles, Monmouth and Ocean counties, New Jersey. N. J. Geological Survey Open-File Map OFM 65, scale 1:24,000.

Wolfe, J. A., 1976, Stratigraphic distribution of some pollen types from the Campanian and lower Maestrichtian rock (Upper Cretaceous) of the middle Atlantic states. U. S. Geological Survey Professional Paper 977, 18 p.

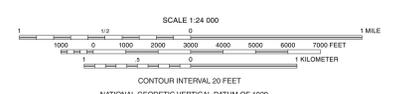
Woolman, Lewis, 1895, Report on artesian wells in southern New Jersey. N. J. Geological Survey Annual Report of the State Geologist for 1894, p. 61-95.



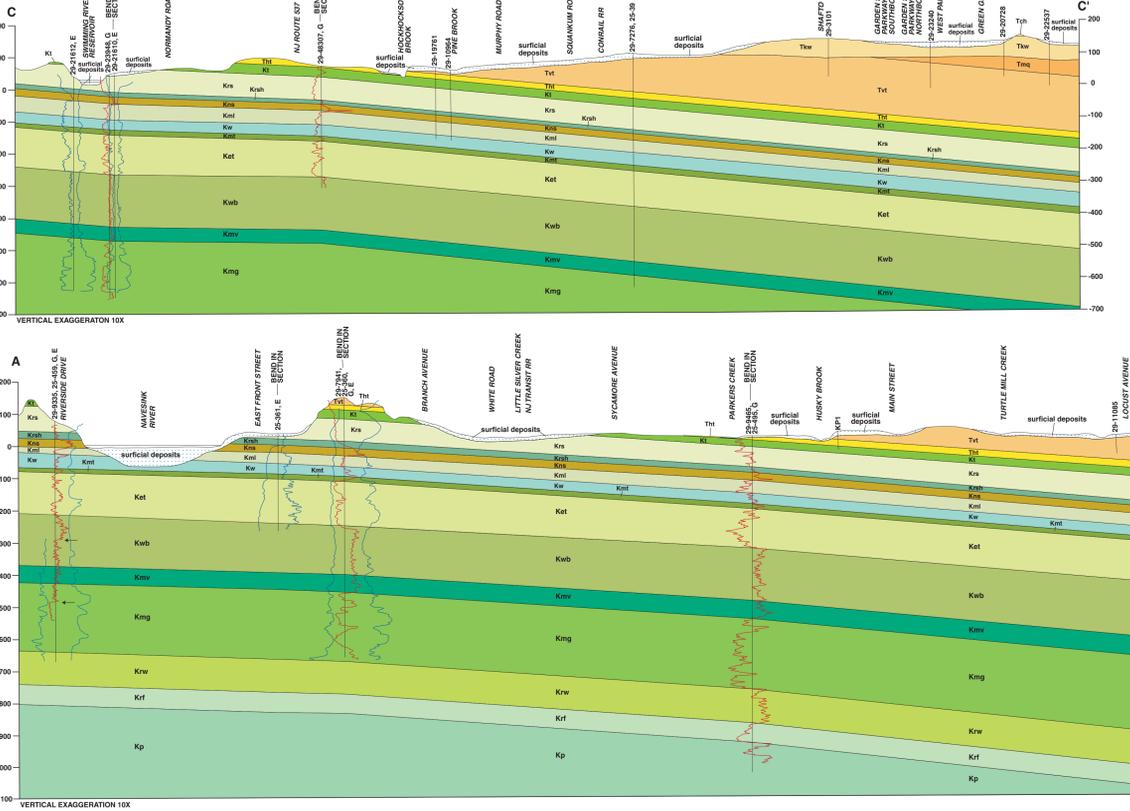
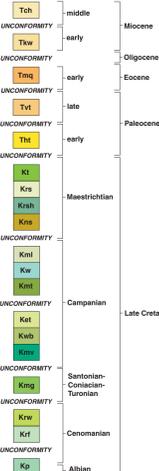
Base map from U. S. Geological Survey, 1954. Photorevised 1981, 1997 North American Datum.

BEDROCK GEOLOGY OF THE LONG BRANCH QUADRANGLE
MONMOUTH COUNTY, NEW JERSEY

by
Scott D. Stanford and Peter J. Sugarmann
2010



CORRELATION OF MAP UNITS



Geology mapped 1994, 2009. Cartography by S. Stanford.
Research supported by the U. S. Geological Survey, National Cooperative Geologic Mapping Program, under U.S. award number 05HQ0008. The views and conclusions contained in this document are those of the authors and should not be interpreted as necessarily representing the official policies, either expressed or implied, of the U. S. Government.