BEDROCK GEOLOGIC MAP OF THE NEW JERSEY PART OF THE

Piedmont bedrock geology mapped by R.A. Volkert, 2012

Retired, New Jersey Geological and Water Survey

Research supported by the U. S. Geological Survey,

National Cooperative Geologic Mapping Progran

under USGS award number G12AC20227. The

views and conclusions contained in this document

as necessarily representing the official policies,

are those of the authors and should not be interpreted

either expressed or implied, of the U. S. Government.

Digital cartography by R.S. Pristas

Coastal Plain bedrock geology mapped by S.D. Stanford, 2012

INTRODUCTION

Bedrock in the New Jersey parts of the Trenton East and Trenton West quadrangles is present in two physiographic provinces, the Piedmont and the New Jersey Coastal Plain. he Fall Line forms the contact between bedrock of the Piedmont and the more easily eroded sediments of the Coastal Plain to the south. Coastal Plain formations are unconsolidated sediments of Upper Cretaceous age. Rocks of the Piedmont extend through the northwestern part of the map area, where they are divided into unmetamorphosed sedimentary rocks of Mesozoic age of the Newark Basin and metamorphosed pre-Mesozoic crystalline rocks of the Trenton Prong. The latter are divisible into northern, central, and southern structural blocks. The Huntingdon Valley fault forms the boundary between the northern and central blocks and the Morrisville fault is the boundary between the central and southern blocks.

Pre-Mesozoic bedrock crops out mainly along the Delaware River, in tributary streams such as Assunpink Creek, and in other minor drainages. It is also encountered in excavations, and in boreholes drilled for environmental investigations or for water wells. Geologic relationships of the pre-Mesozoic rocks have remained poorly understood, in part because of the poor bedrock exposure, and also because of the complexity of the rocks due to their heterogeneity and evidence for multiple episodes of deformation. Early studies were by Bascom and others (1909) and more recent studies on the pre-Mesozoic rocks include that of Owens and Minard (1964). Owens and others (1998), and Monteverde and others (2012). However, these maps provide no information on the age or lithologic characteristics of bedrock, the structural relationships, or joint and fracture data, and the bedrock units are lumped under generic names. The current bedrock mapping provides new information and interpretations on the lithology, stratigraphy, structure, and tectonic framework of the pre-Mesozoic rocks. The mapping of Cretaceous sediments in Owens and Minard (1964) is also updated here. The updates include 1) refinement of contacts, based on additional fieldwork the addition of cross sections, based on logs of 395 water wells and test borings selected from N. J. Geological and Water Survey files (Table 1), and 3) revision of stratigraphic nomenclature and correlations, based on palynologic and lithologic studies completed in the

The Cretaceous formations, and pre-Cretaceous bedrock, are overlain in most of the map area by fluvial and estuarine deposits of Pliocene and Quaternary age. These deposits are shown and described on a separate surficial geologic map (Stanford, 2014).

Coastal Plain since the 1970s.

Cretaceous Sediments

Cretaceous sediments overlie pre-Mesozoic bedrock southeast of the Assunpink Creek valley. They dip to the southeast at between 30 and 50 feet per mile and thicken from a feather edge at their northwest limit to about 800 feet thick in the southeast corner of the renton East quadrangle. The sediments include sand, silt, clay, and glauconite laid down in fluvial, nearshore marine, and continental shelf settings between 95 and 75 Ma (million

The basal Cretaceous unit is the Potomac Formation, which unconformably overlies weathered pre-Mesozoic metamorphic bedrock. The Potomac consists of interbedded quartz sand and clay deposited in fluvial and, possibly, shallow marine settings. The clays were subaerially exposed in floodplains during their deposition and so are commonly oxidized to white, yellow, and red colors (fig. 1). While Owens and Minard (1964) mapped these deposits as the Raritan Formation, they are now known from palynologic studies to be part of the older Potomac Formation. The Raritan overlies the Potomac northeast of the Trenton area but does not extend to the Trenton East and West quadrangles (Owens and others, 1998). Thick beds of coarse sand in the Potomac are productive aguifers and are tapped by

several public supply wells in the map area (wells 136, 150, 154, 192, 264-267, 318, and 319

in Table 1), and numerous domestic wells. The Potomac is unconformably overlain by the Magothy Formation, which consists of lignitic micaceous quartz sand silt and clay commonly interlaminated or in thin interbeds fig. 2). The Magothy was laid down in tidally influenced shallow-shelf and estuarine settings. The uppermost 10 to 30 feet of the Magothy is generally more silty and clayey than the rest of the formation, and is distinguished from the overlying Merchantville Formation by the absence of glauconite. Owens and Minard (1964) mapped this upper silty-clayey section as part of the Merchantville; thus, their Merchantville is thicker, and their Magothy is thinner. than shown here. Sand beds in the Magothy yield water to numerous domestic wells. particularly downdip from the outcrop belt where wells were drilled through the overlying

layey Merchantville and Woodbury formations. The Magothy is unconformably overlain by the Merchantville Formation, which is a glauconitic clayey silt deposited in mid-shelf settings (fig. 3). The Merchantville is the basal unit of a shallowing-upward marine sequence that includes the overlying Woodbury and on-glauconitic clay deposited in prodelta to inner-shelf environments. The Merchantville-Woodbury contact is placed at the highest occurrence of glauconite. The Woodbury is about 30 feet thicker on this map than on Owens and Minard (1964) based on this criterion and on field observations in the Crosswicks-Edgebrook area. The Woodbury transitions upsection to the Englishtown Formation, which is predominantly quartz sand with some interbeds of silt and clay deposited in delta-front to inner-shelf environments. The Woodbury-Englishtown contact is placed at the first occurrence of sand. The Merchantville and Woodbury formations

domestic wells but is too thin and shallow in the map area for extensive use as an aquifer. The northern part of the map area is underlain by Mesozoic rocks that were deposited in the Newark Basin, a northeast-trending rift basin that contains interbedded sedimentary and igneous rocks. These consist of Upper Triassic to Lower Jurassic conglomerate, sandstone, iltstone, and shale of fluvial and lacustrine origin, three tholeiitic basalt formations, and intrusive bodies of diabase. Only the lowest part of this succession, the sedimentary Stockton

Formation, crops out in the map area.

are of low permeability and form a confining unit. The Englishtown yields water to a few

The Upper Triassic Stockton Formation was deposited directly on the eroded surface of the pre-Mesozoic crystalline rocks but nowhere is the contact between them presently exposed. Much of the lower part of the Stockton is arkosic sandstone that contains 50 to 70 percent quartz and 15 to 40 percent feldspar (Van Houten, 1980), which was derived from the underlying crystalline basement. Core samples from a depth of 20 to 27 ft. below ground surface from a monitor well drilled in 1994 along Spruce Street, in Ewing Township (not shown on map), penetrated Stockton Formation containing clasts as much as one inch long of Cambrian-age Chickies Quartzite.

Rocks that are here correlated with the Wissahickon Formation do not crop out in the map area, but are present beneath a cover of Cretaceous to recent sediments and extend as far south as Gloucester and Salem Counties (Volkert and others, 1996). They include mainly schist and gneiss and minor amounts of interlayered amphibolite. Drill core samples of Wissahickon schist recovered beneath the New Jersey Coastal Plain in Gloucester County contain plagioclase + quartz + biotite ± garnet ± sillimanite. They are spatially associated with amphibolites composed of plagioclase + hornblende ± biotite ± garnet. Mineral assemblages of both lithologies are consistent with metamorphism to amphibolite-facies conditions (Volkert and others, 1996). Recent geochronology of felsic metavolcanic rocks in the Wissahickon Formation in the Piedmont of Pennsylvania and Delaware yields a crystallization age of 481 Ma, indicating that the Wissahickon Formation may be as young as Ordovician in age Bosbyshell and others, 2001; Aleinikoff and others, 2006).

Chickies Quartzite crops out in the northwestern part of the Trenton Prong where it is in unconformable contact with Mesoproterozoic rocks. Lithologies representative of the Chickies include: quartz-sericite schist with disseminated, fine grains of black tourmaline; conglomerate and pebbly sandstone; and thin-bedded quartzite (fig. 4). At Morrisville, ennsylvania, we estimate a thickness of about 730 feet for the Chickies. Outcrops of Chickies extend east from Morrisville part way across the Delaware River, but in New Jersey the unit is covered by overburden. However, it is known to be present from drill-hole data.

The Chickies Quartzite was formed from a sedimentary protolith. Locally, the unit displays graded beds and tabular cross-beds that indicate bedding is right-side-up and further substantiates a sedimentary origin. The presence of *Skolithus linearis* burrows in the Chickies suggest that it is Cambrian in age (Bascom and others, 1909). If correct, then it may correlate with the Lower Cambrian Hardyston Quartzite in the New Jersey Highlands and eading Prong in southeastern Pennsylvania.

Neoproterozoic Rocks The Mesoproterozoic rocks locally are intruded by thin, discordant diabase dikes interpreted to be Neoproterozoic. Dikes are hard, dense, unfoliated, and are inmetamorphosed. They are poorly exposed and known mainly through recovery of drill core or rotary cuttings. One dike was encountered in an excavation in the Trenton East in excavations, and in core from a well drilled at Mercer County Park (Monteverde and others, 2012). The dikes are confined to areas underlain by Mesoproterozoic rocks and do not intrude the Chickies Quartzite, Wissahickon schist or gneiss, or Mesozoic rocks. Trenton rong dikes have geochemical compositions that are similar to alkalic basalt and that closely overlap the compositions of dikes in the New Jersey Highlands (Volkert, 2004).

Mesoproterozoic Rocks Mesoproterozoic rocks consist of a heterogeneous assemblage of gneisses that contain the mineral assemblage hornblende + clinopyroxene + plagioclase + quartz ± orthopyroxene + garnet in rocks of felsic to intermediate composition, and hornblende + clinopyroxene plagioclase in mafic rocks, that are representative of metamorphism to granulite-facies conditions. All lithologies are foliated and/or layered. They are divisible into distinct sequences based on their mineralogical and geochemical compositions. These include: 1) thin sheets of the newly-named Assunpink Creek Granite; 2) supracrustal rocks formed from sedimentary and igneous protoliths: and 3) orthogneisses formed from igneous protoliths of felsic, intermediate and mafic composition of the newly-named Colonial Lake Suite. Mesoproterozoic rocks are locally intruded discordantly by thin bodies of unfoliated granite pegmatite composed of quartz + K-feldspar + plagioclase (fig. 6) too small to be shown on the map. Similar pegmatites in the New Jersey Highlands are dated at 1004 to 986 Ma (Volkert

and others, 2005) consistent with emplacement well after the metamorphic peak.

syenogranite on a classification diagram of plutonic rocks (not shown).

The Assunpink Creek Granite is medium to coarse grained, pinkish gray to pinkish white, and weakly foliated to unfoliated potassic granitic rock (fig. 7). It is composed mainly of quartz, K-feldspar, plagioclase, and biotite. Augen of K-feldspar as much as several inches long are common where the rock has been deformed. Assunpink Creek Granite has generally conformable contacts with the gneisses, but locally displays slightly discordant or embayed contacts into the gneisses, thus indicating that it is younger. The granite falls in the field of

Supracrustal rocks include lithologies formed from sedimentary protoliths that are

napped as quartzofeldspathic gneisses, calc-silicate gneiss (fig. 8, 9), minor quartzite and

marble. They are spatially associated with mafic gneisses mapped as amphibolite formed from an igneous protolith. Amphibolite is medium-grained and foliated gneiss composed of plagioclase, hornblende and clinopyroxene (fig. 10). The geochemical composition of amphibolite that is interlayered with metasedimentary gneisses has an affinity to tholeiitic basalt formed in an island arc tectonic setting. Quartzofeldspathic gneiss is pinkish-gray medium-grained, foliated and locally layered rock composed of quartz, K-feldspar, plagioclase biotite, and local garnet and graphite. Calc-silicate gneiss is greenish-gray, medium-grained, foliated and locally layered rock composed of plagioclase, clinopyroxene and local quartz, hornblende, and titanite. The geochemical composition of guartzofeldspathic and calcsilicate gneisses is consistent with protoliths that were mainly sandstone or shale. Quartzite is light gray to pinkish-gray, medium-grained, foliated and locally layered rock composed mainly of quartz and varied amounts of K-feldspar and biotite. It is intercalated with garnetbearing quartzofeldspathic gneiss, but occurs in layers too thin to be shown separately on the map. The geochemical composition of guartzite is consistent with a guartz-rich sedimentary protolith. Carbonate rock is not exposed in the map area but was encountered in drill-holes emplaced in 1960 near downtown Trenton. The rock penetrated in three of the borings is lescribed as "limestone" without any accompanying description of the texture or mineralogy. Bedrock drilled in adjacent borings is described as gneiss or granite, suggesting that the

carbonate rock in these borings may be marble. Because of uncertainty in the age or lithic

description of this rock, it is not shown on the map.

The supracrustal rocks are spatially associated with the Colonial Lake Suite, an assemblage of calc-alkaline magmatic arc rocks comprised of tonalite gneiss, diorite gneiss, gabbro gneiss, and amphibolite. Tonalite gneiss is medium grained, well foliated, and composed of quartz, plagioclase, hornblende, clinopyroxene, biotite, and garnet (fig. 11 Diorite gneiss is medium grained, well foliated, and composed of plagioclase, hornblende linopyroxene, orthopyroxene, and biotite. Quartz, garnet and sulfide minerals occur locally. Gabbro gneiss is medium grained, moderately foliated, and composed of plagioclase, hornblende and clinopyroxene Quartz and sulfide minerals occur locally Amphibolite of the Colonial Lake Suite is a medium-grained and foliated gneiss composed of plagioclase. hornblende and clinopyroxene. It differs from supracrustal amphibolite in that it has a tholeiitic geochemical composition and an affinity to basalt formed in a mid-ocean ridge tectonic setting.

STRUCTURE

Bedding in the Stockton Formation strikes uniformly northeast at an average of N55°E Beds dip northwest at 5° to 19° and average 11°. Bedrock exposure in the map area is insufficient to permit a more detailed directional analysis of bedding data.

Crystallization foliation, formed by the parallel alignment of mineral grains in the pre-Mesozoic metamorphic rocks, defines the strike of the bedrock. It is an inherited feature from compressional stresses during high-grade metamorphism that penetratively deformed

the rocks. Foliations strike uniformly northeast and average N65°E (fig. 12). They dip mainly southeast, and very locally northwest, at an average of 74°. The structural concordance of metamorphic foliations in rocks of the northern, central and southern blocks is likely due to tectonic transposition (reorientation) of original Proterozoic foliations during the Paleozoic Taconian Orogeny and possibly the Alleghanian Orogeny as well.

Metamorphic foliations in Mesoproterozoic rocks are folded by northeast-plunging, northwest-overturned isoclinal antiforms and synforms. Mineral lineations are parallel to the axes of minor folds and range from 6° to 49° toward N56°E to N82°E.

Pre-Mesozoic metamorphic rocks are deformed by three main faults that are, north to south, the Huntingdon Valley, Morrisville, and an unnamed fault. All three strike about N60°E and dip steeply southeast at 70 to 75°. The Morrisville fault merges with, or is cut off by, the Huntingdon Valley fault northeast of the map area. The Huntingdon Valley and Morrisville faults are characterized by an early ductile deformation fabric parallel to metamorphic foliations that is overprinted by brittle deformation fabric. South of the Morrisville fault, the Wissahickon Formation was thrust northwestward against Mesoproterozoic rocks along an unnamed fault. Unfortunately, a more precise interpretation of the structural relationships is not possible because neither the fault nor the Wissahickon are exposed in the map area.

Deformed outcrops of Chickies Quartzite along the Huntingdon Valley fault display a slip lineation on bedding surfaces that plunges an average of 62° to S44°E and displays reverse movement sense involving transport to the northwest. Right-lateral strike-slip movement has also been observed on the fault in the map area, as well as from the Pennsylvania Piedmont (e.g., Hill, 1991; Valentino, 1999). The reverse movement likely accompanied the Taconian Orogeny at circa 450 Ma. The right-lateral strike-slip movement may be much younger, and possibly related to the Alleghanian Orogeny, as proposed by Kroll and others (1999), based on 40Ar/39Ar ages of 311 Ma from white mica formed at greenschist-facies metamorphic conditions along the Huntingdon Valley-Pleasant Grove fault in Maryland.

Joints in bedrock units are characteristically planar, moderately well formed, moderately to widely spaced, and moderately to steeply dipping. Surfaces are smooth and, less commonly, slightly irregular. Joints are spaced from 1 foot to tens of feet apart, although those developed in massive rocks are spaced wider, more irregularly formed and less continuous than joints in well-foliated gneisses. Joints in the Stockton Formation form two sets that have an average strike and dip of N15°W 86° southwest and N68°E 87° southeast. Bedrock exposure is insufficient to permit a more detailed directional analysis. The strike of the dominant joint set in pre-Mesozoic rocks ranges from N70°W to N10°W and averages

N31°W (fig. 13). They dip northeast and, less commonly, southwest at an average of 80°. DESCRIPTION OF MAP UNITS

Englishtown Formation (Upper Cretaceous) – Quartz sand, fine- to medium-grained, minor coarse sand, with thin interbeds of clay and silt. Sand is white, yellow, reddish-yellow, and light gray where weathered, gray where unweathered. Silt and clay are light gray to brown where weathered, dark gray to black where unweathered. The Englishtown is as much as 60 feet thick in the map area; full thickness in the adjacent Allentown quadrangle is 90 feet Owens and Minard, 1966). Sand and clay contain some lignite and mica, and trace amounts of glauconite. Late Cretaceous (early Campanian) in age based on pollen (Wolfe, 1976),

strontium stable-isotope ratios, and calcareous nannofossils (Sugarman and others, 2010). Contact with the underlying Woodbury Formation is gradational. Woodbury Formation (Upper Cretaceous) - Clay, with minor thin beds of very fine quartz sand. Dark gray and black where unweathered, yellowish-brown to light gray where weathered. As much as 80 feet thick. Clay is micaceous and sparsely lignitic. Widely spaced, subvertical joints are present in places. Late Cretaceous (early Campanian) in age based on pollen (Wolfe, 1976) and calcareous nannofossils (Sugarman and others, 2010). Contact with the underlying Merchantville Formation is gradational. In wells, lower contact is placed

at drillers' report of change from gray or black clay in the Woodbury to green clay or marl in Merchantville Formation (Upper Cretaceous) – Glauconitic, micaceous clayey silt to silty fine sand with some interbedded non-glauconitic, micaceous clayey silt. Olive, dark gray, black where unweathered, olive-brown, yellowish-brown, reddish-brown where weathered. size. Iron cementation is common. Widely spaced, subvertical joints are present in places. Late Cretaceous (Santonian-early Campanian) in age based on calcareous nannofossils sugarman and others, 2010). Unconformably overlies the Magothy Formation. In wells,

lower contact is placed at drillers' report of change from green clay or marl in the Merchantville

to gray clay and white or gray sand in the Magothy. On geophysical well logs, lower contact is marked by decreased gamma-ray intensity in the Magothy. Magothy Formation (Upper Cretaceous) – Quartz sand, very fine-, fine-, and mediumgrained, and clay and silt, thin bedded. Sand is white, yellow, light gray where weathered. gray where unweathered. Sand is plane-bedded to low-angle cross-bedded in places. and beds are typically 3 to 6 inches thick, but may be as much as 2 feet thick. Clay and silt are white, yellow, brown, rarely reddish-yellow, where weathered, gray to black where unweathered. Clay-silt beds are typically less than 3 inches thick but may be as much as 1 foot thick. As much as 90 feet thick. Fine-to-medium sands include some lignite and minor mica. Very fine sand and silt beds include abundant mica and lignite, and minor pyrite. Late Cretaceous (Turonian-Coniacian) in age based on pollen (Christopher, 1979; Sugarman and others, 2010). Pollen from the lower part of the Magothy near Mile Hollow is within Zone Va (L. A. Sirkin, written communication, 1990) indicating a late Turonian age (Christopher, 1979). Unit unconformably overlies the Potomac Formation. In wells, lower contact is placed at change from white and gray clay and sand in the Magothy to red and white clay and yellow.

marked by increased gamma-ray intensity in the Potomac. Potomac Formation (Upper Cretaceous) – Quartz sand, fine- to very coarse-grained, locally gravelly, and clay and silt, thick- to thin-bedded. Sand is white, yellow, red, brown, or pink where weathered and light gray where unweathered. Clay and silt are white, red, yellow, brown, pink where weathered and gray to black where unweathered. Weathered beds are more common in the Potomac than in the overlying Magothy Formation. As much as 600 feet thick in the southeast corner of the map area. Medium-to-coarse sand includes some lignite and minor mica in places. Very fine sand and silt beds, and some unweathered clays, include abundant mica and, in places, lignite. The Potomac Formation in the map area is equivalent to the Potomac Formation, unit 3 (Doyle and Robbins, 1977), based on pollen Owens and others, 1998) and is of early Cenomanian age. Unconformably overlies pre-Mesozoic metamorphic bedrock. In wells, contact with bedrock is placed at change from red, yellow, or white clay and sand in the Potomac to saprolite, decomposed bedrock, weathered

brown, and white coarse sand in the Potomac. On geophysical well logs, lower contact is

NEWARK BASIN Stockton Formation (Upper Triassic) – Interbedded sequence of yellowish- to pinkish-gray, rayish-brown, or reddish-brown, medium- to coarse-grained, thin- to thick-bedded, planar to trough cross-bedded, poorly sorted to clast-imbricated conglomerate, ripple cross-laminated arkosic sandstone, reddish-brown, clayey, fine-grained, locally bioturbated sandstone, and siltstone and mudstone in fining-upward sequences. Conglomerate contains pebble-to cobble-sized clasts of subround to round quartz in matrix of coarse sand composed of quartz and feldspar. Basal part of unit is locally deeply weathered. Lower contact is an erosional unconformity. Thickness of unit is approximately 4,000 feet.

TRENTON PRONG Wissahickon schist and gneiss (Lower Ordovician?) – Schist is medium- to dark-gray, medium grained, well layered and foliated, and composed of oligoclase, biotite, K-feldspar, quartz, muscovite, and garnet. Gneiss is tan or pinkish-white, medium grained, foliated, and composed of quartz, K-feldspar, oligoclase, biotite, and muscovite. Schistose rocks differ from gneiss in having a higher content of biotite and less quartz. Unit is not exposed in the

Chickies Quartzite (Lower Cambrian?) – Interbedded sequence of medium-grained, light gray to light greenish-gray quartz-sericite schist with disseminated, fine grains of black tourmaline, less abundant conglomerate composed of pebble-size clasts of blue quartz, and white to light gray, vitreous, thin-bedded quartzite. Locally displays graded beds and trough cross-laminated beds. Contains the trace fossil Skolithus linearis. Thickness is

Diabase dikes (Neoproterozoic) - Dark greenish-gray, unfoliated, dense, hard, medium-Composed of oligoclase to andesine, altered clinopyroxene and small blebs of sulfide. Hornblende is present locally. Unit does not crop out in the map area but is known from excavations and subsurface borings.

o coarse-grained, mainly unfoliated rock composed of quartz, K-feldspar, plagioclase, and biotite. Locally contains augen of K-feldspar as much as several inches long. Unit has generally conformable contacts with adjacent gneisses, but locally preserves slightly discordant and embayed contacts into the gneiss.

Assunpink Creek Granite (Mesoproterozoic) - Pinkish-gray to pinkish-white, medium-

Supracrustal Rocks

grained, foliated gneiss composed of blue or clear quartz, K-feldspar, oligoclase, biotite, and

Quartzofeldspathic gneiss (Mesoproterozoic) – Tan, pinkish-white, or light gray, medium-

local garnet. Graphite is present in trace amounts. Unit includes thin, conformable layers of Ygc Calc-silicate gneiss (Mesoproterozoic) – Light gray to light greenish-gray, medium-grained,

foliated gneiss composed of oligoclase, clinopyroxene, hornblende, and biotite. Quartz and Amphibolite (Mesoproterozoic) - Grayish-black, medium-grained, foliated rock spatially associated with metasedimentary gneisses in the central block. Unit is composed of plagioclase (oligoclase to andesine), hornblende, clinopyroxene, and biotite. Accessory

Magmatic Arc Rocks

minerals include magnetite, epidote and sulfide.

amphibole in a matrix of plagioclase grains.

Colonial Lake Suite Tonalite gneiss (Mesoproterozoic) – Tan, white, or light-gray weathering, light greenish-gray, medium-grained, foliated gneiss composed of blue or clear quartz, oligoclase to andesine,

clinopyroxene, hornblende, biotite, and garnet. Diorite gneiss (Mesoproterozoic) - Light-gray- or tan-weathering, greenish-gray or greenishbrown, medium-grained, foliated rock containing andesine or oligoclase, augite, hornblende, and local orthopyroxene. Quartz, biotite, garnet, and magnetite occur locally. May contain conformable layers of amphibolite or tonalite gneiss.

Gabbro gneiss (Mesoproterozoic) – Black, medium-grained, foliated, dense, rock composed of plagioclase, clinopyroxene, hornblende, and biotite. Unit is spatially associated with diorite Ycla Amphibolite (Mesoproterozoic) – Grayish-black, medium-grained, well foliated rock spatially

associated with felsic and intermediate orthogneisses in the northern block. Composed of

plagioclase (oligoclase to andesine), hornblende, and clinopyroxene. Accessory minerals

include magnetite, epidote and sulfide. Very locally contains abundant small phenocrysts of

Yu Crystalline rocks, undivided (Mesoproterozoic) – Shown in cross section only.

REFERENCES Aleinikoff, J.N., Schenck, W.S., Plank, M.O., Srogi, L.A., Fanning, M.C., Kamo, S.L., and Bosbyshell, H., 2006, Deciphering igneous and metamorphic events in high-grade rocks of the Wilmington Complex, Delaware: Morphology, cathodoluminescence

and backscattered electron zoning, and SHRIMP U-Pb geochronology of zircon and monazite: Geological Society of America Bulletin, v. 118, p. 39-64. Bascom, F., Darton, N.H., Kümmel, H.B., Clark, W.B., Miller, B.L., and Salisbury, R.D., 1909, Description of the Trenton quadrangle, N.J.-Pa: U.S. Geological Survey Geologic

Atlas, Folio 167, 24 p. Bosbyshell, H., Srogi, L., Aleinikoff, J.N., Plank, M.O., Schenck, W.S., Crawford, M.L., and Williams, M., 2001, Tectonic synthesis of recent geochronological and geochemical data in the central Appalachian Piedmont of SE PA and N DEL: Geological Society of America Abstracts with Programs, v. 33, p. A261-A262.

Christopher, R.A., 1979, Normapolles and triporate pollen assemblages from the Raritan and Magothy formations (Upper Cretaceous) of New Jersey: Palynology, v. 3, p.

Doyle, J.A., and Robbins, E.I., 1977, Angiosperm pollen zonation of the Cretaceous of the Atlantic Coastal Plain and its application to deep wells in the Salisbury embayment: Palynology, v. 1, p. 43-78.

Hill, M.L., 1991, Post-Taconic transpression in the Pennsylvania Piedmont, in Crawford, M.L., and Crawford, W.A., eds., Evolution and Assembly of the Pennsylvania-Delaware Piedmont: Field Guide and Proceedings of the 8th Annual Meeting of the

Kroll, M.A., Muller, P.D., and Idleman, B.D., 1999. Late Paleozoic deformation within the Pleasant Grove shear zone, Maryland: Results from 40Ar/39Ar dating of white mica, in Valentino, D.W., and Gates, A.E., eds., The Mid-Atlantic Piedmont: Tectonic missing link of the Appalachians: Geological Society of America Special Paper

map of the Princeton quadrangle, Mercer and Middlesex Counties, New Jersey: New Jersey Geological Survey Open-File Map OFM 93, scale 1:24,000. Owens, J.P., and Minard, J.P., 1964, Pre-Quaternary geology of the Trenton East quadrangle, New Jersey-Pennsylvania: U.S. Geological Survey Geologic Quadrangle Map

New Jersey: U.S. Geological Survey Geologic Quadrangle Map GQ-566, scale

A.A., Jr., and Orndorff, R.C., 1998, Bedrock geologic map of central and southern New Jersey: U.S. Geological Survey Miscellaneous Investigations Map I-2540-B, Stanford, S.D., 2014, Surficial geology of the Trenton East and Trenton West quadrangles, Burlington and Mercer counties, New Jersey: New Jersey Geological and Water

Survey Open-File Map OFM 102, scale 1:24,000.

Sugarman, P.J., Miller, K.G., Browning, J.V., Aubry, M.D., Brenner, G.J., Bukry, D., Butari, B., Feigenson, M.D., Kulpecz, A.A., McLaughlin, P.P., Jr., Mizintseva, S., Monteverde, D.H., Olsson, R., Pusz, A.E., Rancan, H., Tomlinson, J., Uptegrove, J., and Velez, C.C., 2010, Medford site, in Miller, K.G., Sugarman, P.J., and Browning, J.V., eds., Proceedings of the Ocean Drilling Program, Initial Reports, v. 174AX (Supplement): College Station, Texas, Ocean Drilling Program, p. 1-93.

Valentino, D.W., 1999, Late Paleozoic dextral transpression in the crystalline core of the Pennsylvania reentrant, in Valentino, D.W., and Gates, A.E., eds., The Mid-Atlantic Piedmont: Tectonic missing link of the Appalachians: Geological Society of America Special Paper 330, p. 59-71.

Van Houten, F.B., 1980, Late Triassic part of Newark Supergroup, Delaware River section, west-central New Jersey, in Manspeizer, W., ed., Field Studies of New Jersey Geology and guide to field trips: Rutgers University Press, p. 264-276. Volkert, R.A., 2004, Geochemistry and tectonic setting of Late Neoproterozoic diabase dikes, New Jersey Highlands and Trenton Prong, in Puffer, J.H., and Volkert, R.A., eds.,

Field Guide and Proceedings of the 21st Annual Meeting of the Geological Association of New Jersey, p. 27-51. Volkert, R.A., Drake, A.A., Jr., and Sugarman, P.J., 1996, Geology, geochemistry, and tectonostratigraphic relations of the crystalline basement beneath the Coastal Plain

of New Jersey and contiguous areas: U.S. Geological Survey Professional Paper Volkert, R.A., Zartman, R.E., and Moore, P.B., 2005, U-Pb zircon geochronology of Meso-

proterozoic postorogenic rocks and implications for post-Ottawan magmatism and metallogenesis, New Jersey Highlands and contiguous areas, USA: Precambrian Wolfe, J.A., 1976, Stratigraphic distribution of some pollen types from the Campanian and

Figure 1. Interbedded clay and sand of the Potomac Formation (below shovel

head) overlain by surficial deposit (above shovel head).

lower Maestrichtian rocks (Upper Cretaceous) of the Middle Atlantic states: U.S.





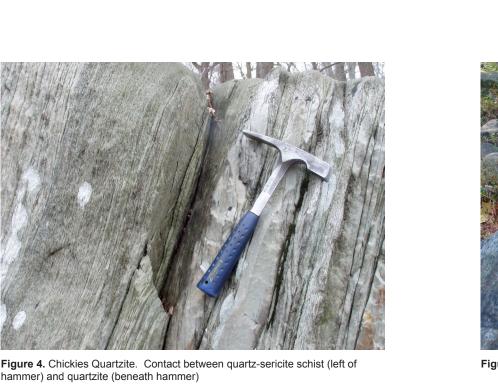
Figure 7. Assunpink Creek Granite



gray and white beds) of the Magothy Formation.



to reddish-brown sediment (above shovel head) is weathered and iron-cemented. Dark grav sediment (below shovel head) is unweathered. Yellowish-brown sediment below shovel is slumped from above.



surficial deposits





Figure 12. Rose diagram showing the orientation of crystallization folliation in pre-Mesozoic rocks.

Sector angle = 10°

Sector angle = 10° Figure 13. Rose diagram showing the orientation of joints in



gray and more deeply weathered.



quartzofeldspathic gneiss (right of hammer).



LINEAR FEATURES

Boring – Rock type at bottom: M, Mesoproterozoic rocks; W, Wissahickon gneiss or schist; C, Chickies

Well or boring with log in Table 1 – Location accurate to within 200 feet. Formations and lithologies based

²¹⁶ Well or boring with log in Table 1 – Location accurate to within 500 feet. Formations and lithologies based

Cretaceous formation observed in former excavation or exposure – From Owens and Minard (1964)

OKet ■ Cretaceous formation penetrated in power-auger hole – From Owens and Minard (1964). Where multiple

Zone Va * Pollen zone – Symbol indicates location of sample in Magothy Formation. Pollen analysis by L. A. Sirkin

Gamma-ray log – On Cretaceous cross section C-C'. Gamma-ray intensity increases to right

— — Form line showing trend of bedding and foliation in pre-Cretaceous rocks in cross section

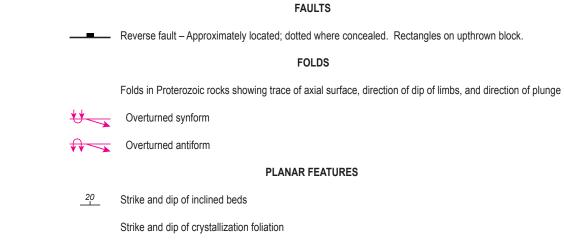
formations were penetrated, number preceding formation abbreviation is depth, in feet below land surface, of

Cretaceous formation observed in exposure, excavation, or hand-auger hole

base of formation. Surficial deposits indicated by "s".

Contact – Approximately located; dashed where concealed by water. Triangle indicates contact observed in

EXPLANATION OF MAP SYMBOLS



Bearing and plunge of mineral lineation

Quartzite; S, Stockton Fm.

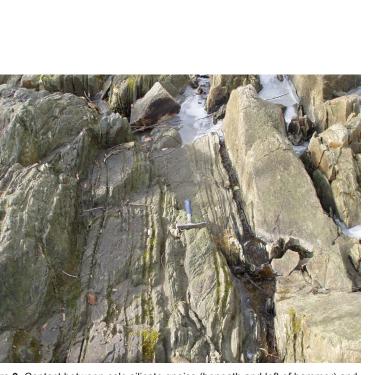
on drillers' or loggers' descriptions.

(written communication, 1990).

ത്ര Location of photo

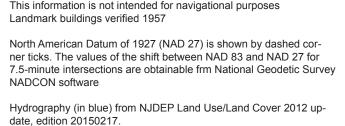
Inclined

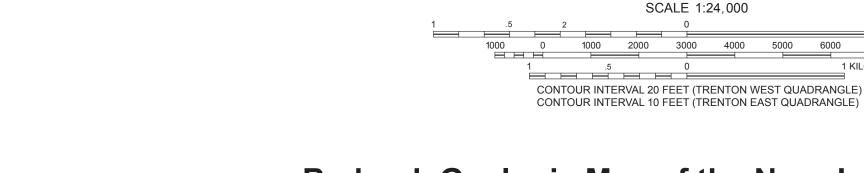
→ Vertical





Trenton East topography compiled 1942. Planimetry derived from imagery taken 1995. Survey control current as of 1943. Trenton West topography compiled 1947. Planimetry derived from imagery taken 1995. Survey control current as of 1957. Selected hydrographic data compiled from NOS Chart 296 (1954)





CORRELATION OF MAP UNITS

Unconformity

Unconformity

Unconformity

PIEDMONT

Newark Basin

Unconformity

Trenton Prong

Fault Contact

€c

Unconformity

Zd

Intrusive Contact

Granitic Rocks

Assunpink Creek Granite

Intrusive Contact

Supracrustal Rocks

Magmatic Arc Rocks

Colonial Lake Suite

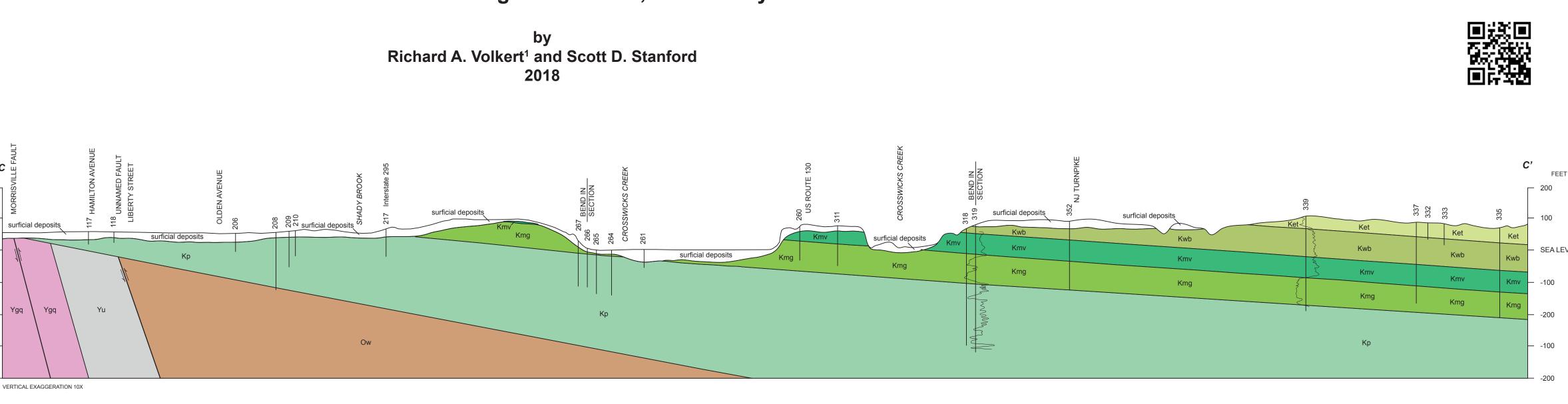
CRETACEOUS

ORDOVICIAN

NEOPROTEROZOIC

MESOPROTEROZOIC

Bedrock Geologic Map of the New Jersey Part of the **Trenton West and Trenton East Quadrangles** Mercer and Burlington Counties, New Jersey



KORDBNTOWN

NO VERTICAL EXAGGERATION

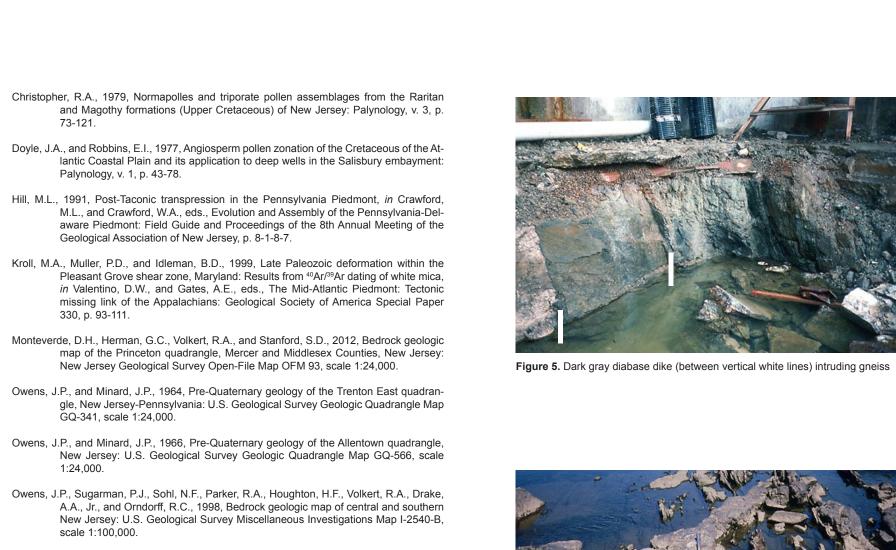




Figure 6. Course-grained pegmatite (beneath hammer) cutting across outcrops of quartz-feldspathic gneiss. Pegamtite extends from lower left corner of photo to upper right. Gneiss runs from top to bottom of photo.





Figure 8. Supracrustal quartzofeldspathic gneiss. Biotite-rich layers are dark







surficial deposits

VERTICAL EXAGGERATION 10X

SEA LEVEL

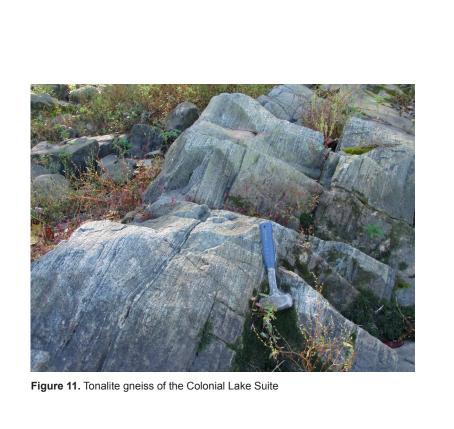


Figure 10. Supracrustal amphibolite. Keys for scale.

NADCON software

Bedrock Geologic Map of the New Jersey Part of the Trenton West and Trenton East Quadrangles Mercer and Burlington Counties, New Jersey

New Jersey Geological and Water Survey Open File Map 122 2018

Pamphlet with table 1 to accompany map

Table 1. Selected well and boring records. Footnotes at end of table (p. 8)

| Well Number | Identifier ¹ | Formations Penetrated ² |
|----------------|-------------------------|------------------------------------|
| 1 | 27-1664 | 34 Qwf 100 ss&sh |
| 2 | 27-5896 | 33 Qwf, cased to 33, yield 25 gpm |
| 3 | 27-2918 | 5 Qwf, cased to 55, yield 25 gpm |
| 4 | 27-2751 | 40 Qwf 90 ss |
| 5 | 27-2288 | 15 Tp 402 ss |
| 6 | 27-8783 | 8 Qwf 25 wss |
| 7 | 27-10012 | 15 Qwf 37 wss 65 ss&sh |
| 8 | 28-3500 | 18 Tp 22 ss |
| 9 | 28-5471 | 12 Tp 70 ss |
| 10 | 27-331 | 25 Tp+wss 41 ss 70 sh |
| 11 | 27-2550 | 15 Qal+wss 75 ss |
| 12 | 28-28045 | 19 wss |
| 13 | 27-2587 | 62 wss 200 ss |
| 14 | 28-94 | 40 wss 201 ss |
| 15 | 27-2747 | 28 wss 96 ss 337 gn |
| 16 | 28-16103 | 50 wss 170 ss 290 gn |
| 17 | 28-493 | 20 Tp+wss 48 wss 198 ss 265 gn |
| 18 | 28-28456 | 4 Tp 35 wss |
| 19 | 28-29954 | 20 Tp+wss 46 wss |
| 20 | 27-10250 | 20 Tp 29 wsc 31 sc |
| 21 | 27-10230 | 52 wss |
| 22 | 27-11256 | 8 Qwf 12 wsc |
| 23 | 27-11250 | 5 f 21 Qwf 25 gn |
| 24 | 27-11962 | 13 f 33 Qwf 41 gb |
| 25 | 27-10985 | 5 f 13 Qwf 37 wr |
| 26 | 28-29900 | 2 f 15 wsc 20 sc |
| 27 | 27-11068 | 15 f 42 Qwf 43 wr 60 r |
| 28 | 28-31375 | 19 f 29 Qwf 30 wgn 35 gn |
| 29 | 27-8253 | 13 f 22 Qwf 27 gn |
| 30 | DOT 163W-17 | 19 Qwf 29 sc |
| 31 | 28-23426 | 10 f 40 Qwf 41 sc |
| 32 | 28-23520 | 5 f 14 Qwf 24 gn≻ |
| 33 | DOT 379W-8 | 36 Qwf |
| 34 | 28-20864 | 3 f 10 Qwf 17 wr |
| 35 | 28-25242 | 6 f 22 Qwf 42 wsc |
| 36 | 28-22049 | 5 f 14 Qwf 16 wr |
| 37 | DOT 439W-24 | 21 Owf |

| Well | Identifier ¹ | Formations Penetrated ² |
|----------|-----------------------------|--|
| Number | | |
| 38 | 28-23512 | 28 Qwf 30 r |
| 39 | 28-21509 | 9 f 23 Qwf 29 wgn 30 gn |
| 40 | 28-22999 | 4 f 32 Qwf 43 wgn 50 gn |
| 41 | 28-17885 | 6 f 16 Qwf 50 gn&cg |
| 42 | 28-22175 | 35 f+Qwf 60 wsc≻ (average of numerous borings in small area) |
| 43 | 28-30064 | 18 f 34 Qwf 60 wr |
| 44 | DOT L94A | 9 f 25 Qwf 40 sc |
| 45 | 28-25963 | 42 Qwf 43 wr? |
| 46 | DOT 379W-17 | 23 Qwf 25 wgn 30 gn |
| 47 | 28-23692 | 38 Qwf 39 sc |
| 48 | DOT 125W-194 | 46 Qwf |
| 49 | 28-28263 | 45 Qwf |
| 50 | DOT B1001 | 30 f+Qwf 49 gn 58 q |
| 51 | DOT DR 1 | 5 w 6 Qal 11 gn |
| 52 | DOT DR 3 | 13 w 14 Qal 19 sc |
| 53 | DOT DR 4 | 4 w 22 Qal 26 sc |
| 54 | DOT DR 5 | 13 w 17 Qal 22 gn |
| 55 | 28-19648 | 30 f+Qwf 98 wsc 113 sc |
| 56 | 28-28324 | 40 Qwf |
| 57 | DOT L68 | 15 f 40 Qwf |
| 58 | DOT L31 | 9 f 28 Qwf 39 q |
| 59 | 28-32138 | 7 f 17 Qt 27 Qwf 47 wsc |
| 60 | DOT DR 7 | 3 w 26 Qal 31 q≻ |
| 61 | DOT DR 9 | 8 w 53 Qal |
| 62 | DOT DR 10 | 4 w 24 Qal 30 sc |
| 63 | DOT DR 11 | 8 w 60 Qal |
| 64 | DOT DR 12 | 7 w 53 Qal 60 q |
| 65 | DOT DR 13 | 5 w 38 Qal 60 Kp |
| 66 67 | DOT DR 15 DOT DR 17 | 6 w 28 Qal 63 Kp |
| | | 4 w 17 Qal 60 Kp 5 f 28 Qwf 80 Tp? 200 Kp |
| 68 69 | 28-4078 DOT L53 | |
| 70 | DOT 125W-190 | 11 f 40 Qwf 9 f 45 Qwf 50 Tp? |
| | DOT 125W-190 DOT 125W-77 | 7 f 43 Qwf |
| 71 | | ` |
| 72 73 | DOT 125W-101 | 29 f 41 Qwf 43 Kp |
| 74 | DOT W-12 | 45 Qal 66 Kp 5 w 23 Qal 60 Kp |
| 75 | DOT DR 19 | ` ' |
| | DOT DR 20 | 31 w 50 Qal 60 Kp |
| 76 | DOT DR 22 | 12 w 44 Qal 60 Kp |
| 77 78 | DOT DR 23 DOT DR 25 | 37 w 50 Qal 60 Kp 9 w 46 Qal 59 Kp |
| 79 | | |
| 80 | 28-13021 28-13022 | 13 f 21 Qst 30 Qwf 42 Kp 22 f 41 Qwf 66 Kp |
| 81 | DOT 125W-122 | 32 Qwf 56 Kp |
| 82 | DOT 125W-122 DOT DR 37 | 4 w 28 Qal 60 Kp |
| 82 | 28-1600 | 4 w 28 Qai 60 Kp 50 Tp+wr 78 wr 194 ss |
| 84 | 28-13146 | 50 1p+wr 78 wr 194 ss 4 f 17 Qwf 22 wr? |
| 85 | 28-13146 | 6 f 19 Owf 20 wsc |
| 86 | 28-13787 | 27 Qwf 30 q, screened 20-30, yield 60 gpm |
| 87 | 28-29354 | 9 f 19 Qwf 23 wgn 40 gn |
| 88 | | 30 Qwf, screened 28-30, yield 12 gpm |
| | 28-10291 | |
| 89 | 28-28670 | 6 f 27 wsc |
| 90 | DOT 439W-15 | 9 f 15 wgn 20 gn |
| 91 | DOT 379W-10 | 19 Qwf 29 gn |
| 92 | DOT S9 | 3 f 32 Qwf 43 gn |
| 93 94 | 28-575 | 36 Qwf, screened 26-36, yield 140 gpm |
| 94 | 28-699 | 38 Qwf 39 r, screened 18-38, yield 72 gpm |

| Well Number | Identifier ¹ | Formations Penetrated ² |
|----------------|-------------------------|---|
| 95 | 28-2711 | 37 Qwf 42 r, screened 22-38, yield 50 gpm |
| 96 | 28-27883 | 25 Qwf 40 wgn 87 gn |
| 97 | 28-26479 | 15 Qwf 16 gn |
| 98 | 28-23390 | 48 Qwf 49 r |
| 99 | 28-5502 | 54 Tp 55 wr?, screened 50-53, yield 10 gpm |
| 100 | 28-7173 | 41 Qwf 45 wr |
| 101 | 28-3437 | 24 Qwf 45 Tp? |
| 102 | 28-6961 | 22 Qwf 41 Tp? |
| 103 | 28-4896 | 40 Qwf 70 wsc, logged by F. J. Markewicz 10/13/64, screened 40-46, yield 10 gpm |
| 104 | 28-15574 | 48 Qwf 53 wr |
| 105 | 28-9906 | 46 Qwf 47 wr 50 r, screened 40-50, yield 40 gpm |
| 106 | 28-7910 | 50 Qwf, screened 45-50, yield 28 gpm |
| 107 | 28-6972 | 55 Qwf, screened 47-50, yield 30 gpm |
| 108 | 28-8437 | 52 Qwf 53 r, screened 49-52, yield 25 gpm |
| 109 | 28-9951 | 51 Qwf 55 wr, screened 45-50, yield 60 gpm |
| 110 | DOT S11 | 4 f 56 Qwf 63 Kp |
| 111 | DOT S19 | 33 Tp 61 Kp |
| 112 | DOT S26 | 45 Tp 71 Kp |
| 113 | 28-3578 | 12 Qwf 52 Kp |
| 114 | 28-5562 | 45 Qwf 61 Kp |
| 115 | 28-25965 | 27 Qwf |
| 116 | 28-4480 | 15 f 28 Qwf 42 wr? |
| 117 | 28-15680 | 30 Qwf 33 Kp |
| 118 | 27-1815 | 17 Qwf 35 Kp |
| 119 | 28-1576 | 17 Qwf 30 Kp |
| 120 | 27-1792 | 18 Qwf 30 KP |
| 121 | 28-3694 | 30 Qwf 79 Kp |
| 122 | 28-11247 | 15 Qwf 100 Kp |
| 123 | 28-2356 | 41 Tp 134 Kp |
| 124 | DOT S36 | 3 f 26 Qt 35 Kp |
| 125 | DOT S219 | 31 Qt 51 Kp |
| 126 | 28-5500 | 30 Qt 51 Kp |
| 127 | 28-3667 | 20 Tp 60 Kp |
| 128 | 28-9467 | 30 Tp 56 Kp |
| 129 | 28-6643 | 15 Tp 73 Kp |
| 130 131 | 28-5958 28-3554 | 12 Tp 65 Kp 18 Tp 60 Kp |
| 131 | 28-12801 | 14 Qt 40 Tp 58 Kp |
| 133 | 28-3184 | 6 Qt 53 Kp |
| 134 | 28-3185 | 10 Qt 55 Kp |
| 135 | 28-25629 | 11 Tp 44 Kp |
| 136 | 28-29222 | 11 Tp 138 Kp |
| 137 | 28-8376 | 25 Tp 80 Kp |
| 138 | 28-5559 | 28 Tp 71 Kp |
| 139 | 28-7130 | 8 Tp 71 Kp |
| 140 | 27-4395 | 23 Tp 70 Kp |
| 141 | 28-5986 | 15 Tp 69 Kp |
| 142 | 28-6011 | 39 Tp 71 Kp |
| 143 | 28-5699 | 38 Tp 61 Kp |
| 144 | 28-2788 | 16 Tp 61 Kp |
| 145 | 28-3646 | 18 Tp 71 Kp |
| 146 | 28-6576 | 38 Tp 85 Kp |
| 147 | 28-6418 | 25 Tp 100 Kp |
| 148 | 28-7896 | 16 Tp 90 Kp |
| 149 | 28-1350 | 11 Tp 82 Kp |
| 150 | 28-2927 | 7 Tp 155 Kp 160 wr |
| 151 | DOT 379W-9 | 25 Qwf 33 gn |

| Well | Identifier ¹ | Formations Penetrated ² |
|---------------|-------------------------|---|
| Number 152 | 28-2427 | 30 Tp 89 Kp |
| 153 | 28-5201 | 18 Tp 101 Kp |
| 154 | 28-6625 | 34 Tp 218 Kp 236 wr |
| 155 | 28-161 | 40 Tp 213 Kp 220 Kp or wr |
| 156 | 28-1363 | 40 Tp 215 Kp 235 wsc, logged by F. J. Markewicz |
| 157 | 28-26569 | 32 Tp 55 Kp |
| 158 | 28-2792 | 20 Tp 30 Kmg 230 Kp |
| 159 | 28-4218 | 45 Tp 120 Kp |
| 160 | 28-5306 | 28 Tp 113 Kp |
| 161 | 28-13775 | 6 Qt 105 Kp |
| 162 | 28-13557 | 48 Tp 78 Kp |
| 163 | 28-10701 | 34 Tp 100 Kp |
| 164 | 28-12799 | 45 Tp 56 Kmg 180 Kp |
| 165 | 28-31367 | 25 Tp 65 Kmg 100 Kp |
| 166 | 28-7380 | 10 Tp 65 Kmg 70 Kp |
| 167 | 28-15342 | 50 Kmg 75 Kp |
| 168 | 28-6017 | 7 g 8 Qt 90 Kmg |
| 169 | SC 137 | 8 Qt 128 Kmg |
| 170 | 28-6222 | 10 Qt 62 Kmg |
| 171 | 28-19775 | 15 Qt 113 Kmg 150 Kp |
| 172 | SC 138 | 12 Qt 125 Kmg |
| 173 | SC 139 | 131 Kmg |
| 174 | SC 140 | 6 Qt 128 Kmg |
| 175 | SC 141 | 143 Kmg |
| 176 177 | SC 142 SC 143 | 11 Kmv 85 Kmg 149 Kp |
| 177 | SC 143 SC 144 | 6 g 20 Tp 89 Kmg 20 Tp 93 Kmg 162 Kp |
| 179 | SC 144 SC 145 | 10 Tp 85 Kmg 150 Kp |
| 180 | SC 143 | 16 Tp 108 Kmg 148 Kp |
| 181 | SC 147 | 68 Kmg 134 Kp |
| 182 | 28-31032 | 36 Tp 96 Kmv 160 Kmg |
| 183 | 28-26539 | 12 Tp 50 Kmv 100 Kmg |
| 184 | SC 149 | 9 Qt 134 Kmg |
| 185 | SC 150 | 15 Qt 70 Kmg 131 Kp |
| 186 | 28-7031 | 8 Tp 61 Kmg |
| 187 | SC 151 | 10 Qal 126 Kmg |
| 188 | SC 152 | 75 Kmg 135 Kp |
| 189 | 28-9905 | 31 Tp 43 Kmv 117 Kmg 160 Kp |
| 190 | SC 153 | 28 Tp 90 Kmg 144 Kp |
| 191 | SC 154 | 19 Kmv 79 Kmg 143 Kp |
| 192 | 28-49883 | 151 Kp |
| 193 | 28-7667 | 36 Tp 140 Kp |
| 194 | 28-6241 | 38 Tp 109 Kp |
| 195 | DOT S54 | 4 f 61 Kp |
| 196 | 28-22497 | 18 Tp 52 Kp |
| 197 | 28-5081 | 34 Tp 96 Kp |
| 198 | 28-5010 | 18 Tp 97 Kp |
| 199 | 28-30274 | 31 Tp 45 Kmg |
| 200 | 28-12399 | 55 Tp 68 Kmg 2 f 43 Tp |
| 201 | 28-17124 28-19133 | 2 f 35 Tp |
| 202 | DOT S63 | 8 f 24 Tp 81 Kp |
| 204 | 28-3105 | 16 Tp 92 Kp |
| 205 | 28-30002 | 20 Qwf 59 Kp |
| 206 | 28-29938 | 40 Qwf 70 Kp |
| 207 | 28-1152 | 4 f 8 Tp 60 Kp |
| 208 | 28-1551 | 170 Kp 185 sc, logged by E. S. Lenker |
| | | 1 · · · · · · · · · · · · · · · · · · · |

| Well | Identifier ¹ | Formations Penetrated ² |
|------------|-------------------------|--|
| Number | | |
| 209 | 28-7150 | 2 f 5 Qs 15 Qwf 109 Kp |
| 210 | 28-6061 | 4 f 10 Qwf 68 Kp |
| 211 | 28-30969 | 70 Tp |
| 212 | 28-16493 | 35 Qwf |
| 213 214 | 28-30994 | 5 f 40 Qwf |
| 214 | 28-29315 | 32 Qwf |
| 216 | 28-21-485 28-7129 | 46 Qwf 98 Kp 179 w sc |
| 217 | DOT S69 | 10 Qwf 59 Kp 22 Tp 86 Kp |
| 218 | 28-9204 | 18 Tp 38 Kmv 135 Kmg 170 Kp |
| 219 | 28-11690 | 10 Tp 35 Kmv 135 Kmg 150 Kp |
| 220 | 28-5202 | 28 Tp 40 Kmv 66 Kmg |
| 221 | 28-4964 | 36 Qt 115 Kmg 124 Kp |
| 222 | SC 155 | 16 Tp 65 Kmg 128 Kp |
| 223 | SC 156 | 3 Tp 124 Kmg |
| 224 | SC 157 | 13 Qt 112 Kmg |
| 225 | SC 158 | 19 Qt 112 Kmg |
| 226 | SC 159 | 14 Qt 80 Kmg |
| 227 | SC 160 | 12 Qt 84 Kmg |
| 228 | 28-7663 | 55 Qt 75 Kmg 115 Kp |
| 229 | SC 161 | 17 Qt 69 Kmg 81 Kp |
| 230 | SC 162 | 22 Qt 52 Kmg 78 Kp |
| 231 | SC 163 | 36 Qt 50 Kmg 82 Kp |
| 232 | SC 164 | 39 Qt 75 Kmg 108 Kp |
| 233 | 28-5317 | 18 Qt 63 Kmv+Kmg 117 Kmg 178 Kp |
| 234 | 28-2025 | 14 Qt 90 Kmv+Kmg 105 Kmg |
| 235 | 28-2458 | 16 Qt 23 Kmv 134 Kmg 215 Kp |
| 236 | 28-17765 | 9 Qt 45 Kmy 178 Kmg 200 Kp |
| 237 | 28-10158 | 34 Kwb 93 Kmv 142 Kmg |
| 238 | 28-16886 | 15 Qt 75 Kwb 90 Kmv 150 Kmg |
| 239 | 28-7144 | 14 Qt 81 Kwb+Kmv 143 Kmg |
| 240 | 28-25995 | 11 Qt 88 Kmv 172 Kmg |
| 241 | 28-3824 | 12 Qt 70 Kwb+Kmv 140 Kmv+Kmg 162 Kmg |
| 242 | 28-20542 | 10 Qt 95 Kwb+Kmv 130 Kmg |
| 243 | 28-11057 | 6 Qt 23 Kwb 68 Kmv 145 Kmg |
| 244 | 28-15230 | 8 Qt 45 Kwb 81 Kmv 132 Kmg |
| 245 | 28-7253 | 8 Qt 100 Kwb+Kmv 168 Kmg |
| 246 | 28-10299 | 4 g 11 Qt 85 Kwb+Kmv 91 Kmv 150 Kmg |
| 247 | 28-14224 | 3 f 10 g 25 Qt 115 Kwb+Kmv 230 Kmg 285 Kp |
| 248 | 28-1671 | 10 Tp 178 Kwb+Kmv 207 Kmg |
| 249 | 28-7497 | 5 Tp 155 Kwb+Kmv 213 Kmg 238 Kp |
| 250 | 28-28500 | 60 Kwb 80 Kmv 135 Kmv+Kmg 200 Kmg |
| 251 | 28-19515 | 10 Tp 68 Kwb 105 Kmv 181 Kmg |
| 252 | 28-31081 | 29 Tp 54 Kwb 121 Kmv 208 Kmg |
| 253 | 28-7929 | 28 Qt 96 Kmv 165 Kmg |
| 254 | 28-8677 | 24 Qt 90 Kmv 149 Kmg |
| 255 | 28-4280 | 15 Qt 90 Kmv 119 Kmg |
| 256 | 28-3956 | 20 Qt 90 Kmg, logged by D. G. Parrillo, 11/28/60 |
| 257 | 28-9231 | 4 f 6 Qt 26 Kmv+ Kmg 90 Kmg |
| 258 | 28-1729 | 6 Qt 34 Kmv 108 Kmg |
| 259 | 28-25886 | 15 Qt 49 Kmv 96 Kmg |
| 260 | 28-14445 | 20 Qt 50 Kmv 115 Kmg |
| 261 | SC165 | 16 Qm 35 Kmg 57 Kp |
| 262 | SC 167 | 36 Qm 56 Kmg |
| 263 | SC 168 | 19 Qt 60 Kmg |
| 264 | 28-5409 | 9 Qm 15 Kmg 140 Kp |
| 265 | 28-8769 | 2 f 10 Qm 20 Kmg 137 Kp |

| Well | Identifier ¹ | Formations Penetrated ² |
|---------------|-------------------------|------------------------------------|
| Number 266 | 28-34305 | 12 f over Qm 22 Kmg 122 Kp |
| 267 | 29-5150 | 8 Qt 16 Kmg 127 Kp |
| 268 | 28-31914 | 14 Kmg 53 Kp |
| 269 | 28-104 | 18 Qt 30 Kmg 120 Kp |
| 270 | core 1 (Southgate, | 19 Qm |
| 270 | 2010) | 19 Qili |
| 271 | DOT S88 | 17 Qm 42 Qal 120 Kp |
| 272 | DOT SB341 | 10 Qm 43 Qal 55 Kp |
| 273 | DOT SB119 | 13 Qm 69 Qal 79 Kp |
| 274 | DOT SB50 | 13 Qm 51 Qal |
| 275 | 28-26036 | 30 Qwf 145 Kp |
| 276 | DOT H41 | 3 f 12 Qm 30 Qal |
| 277 | DOT H40 | 6 f 17 Qm 35 Qal |
| 278 | DOT H38 | 9 f 17 Qm 30 Qal |
| 279 | 28-25876 | 21 Qm 28 Qal 60 Kp |
| 280 | DOT 125W-161 | 12 Qm 28 Qal 36 Kp |
| 281 | DOT 125W-76 | 40 Qt 43 Kp |
| 282 | DOT DR 36 | 3 w 17 Qal 59 Kp |
| 283 | DOT DR 37 | 4 w 28 Qal 60 Kp |
| 284 | DOT DR 40 | 7 w 14 Qal 60 Kp |
| 285 | DOT DR 41 | 26 w 36 Qal 60 Kp |
| 286 | 28-9524 | 25 f 63 Qt 73 Kp |
| 287 | 28-9525 | 25 f 63 Qt 73 Kp |
| 288 | DOT DR 55 | 15 w 61 Qal |
| 289 | DOT SB 190 | 18 Qm 40 Qal 66 Kp |
| 290 | DOT SB223 | 21 Qm 40 Qal 61 Kp |
| 291 | DOT SB232 | 17 Qm 40 Qal 66 Kp |
| 292 | DOT H14 | 18 Qm 28 Qal 46 Kp |
| 293 | DOT H20 | 18 Qm 31 Qal |
| 294 | DOT DR 64 | 7 w 50 Qal 60 Kp |
| 295 | DOT DR 68 | 30 w 43 Qal 60 Kp |
| 296 | DOT DR 69 | 8 w 36 Qal 59 Kp |
| 297 | DOT DR 72 | 7 w 30 Qal 60 Kp |
| 298 | DOT DR 73 | 17 w 50 Qal 61 Kp |
| 299 | DOT DR 74 | 25 w 45 Qal 60 Kp |
| 300 | DOT F7 | 40 Qt 64 Kp |
| 301 | SC 172 | 23 Qm 27 Qal 55 Kmg |
| 302 | SC 173 | 32 Qm 47 Qal 55 Kp |
| 303 | SC 174 | 35 Qm 42 Qal 55 Kp |
| 304 | SC 175 SC 176 | 20 Qm 55 Qal |
| 305 306 | SC 176 SC 177 | 33 Qm 55 Qal 12 Qm 18 Qal 55 Kp |
| 306 | SC 177 | 6 f 19 Qt 62 Kp |
| 308 | SC 178 SC 169 | 61 19 Qt 62 Kp 62 Qt 127 Kmg+Kp |
| 309 | SC 170 | 38 Qt 106 Kmg |
| 310 | SC 170 | 20 Qt 85 Kmg 109 Kmg+Kp |
| 311 | 28-9683 | 10 Qt 55 Kmv 123 Kmg |
| 312 | 28-15804 | 10 Qt 55 Kmv 125 Kmg |
| 313 | 28-20500 | 8 Qt 32 Kmv 65 Kmg |
| 314 | 28-4291 | 12 Qt 45 Kmv 75 Kmg |
| 315 | 28-29149 | 32 Tp 115 Kwb+Kmv 180 Kmg |
| 316 | 28-16367 | 40 Tp 155 Kwb+Kmv 172 Kmg |
| 317 | 28-249 | 13 Tp 35 Kwb 75 Kmv 276 Kmg 397 Kp |
| 318 | 28-5042 | 15 Tp 110 Kwb+Kmv 239 Kmg 372 Kp |
| 319 | 28-57004, G | 10 Tp 110 Kwb+Kmv 245 Kmg 397 Kp |
| 320 | 28-21105 | 29 Qt 49 Kwb 94 Kmv 162 Kmg |
| 321 | 28-16829 | 16 Qt 85 Kwb 160 Kmv 205 Kmg |
| U-1 | _ = 5 100=7 | 1 2 |

| Well | Identifier ¹ | Formations Penetrated ² |
|---------------|-------------------------|---|
| Number 322 | 28-17754 | 15 Ot 65 Vish 115 Vmv 145 Vmv 145 Vmv 1400 Vma |
| 323 | 28-15414 | 15 Qt 65 Kwb 115 Kmv 145 Kmv+Kmg 200 Kmg 25 Qt 125 Kwb+Kmv 140 Kmv 187 Kmg |
| 324 | 28-8832 | 18 Qt 75 Kwb 95 Kmv |
| 325 | 28-16759 | 16 Qt 80 Kwb 140 Kmv 193 Kmg |
| 326 | 28-21288 | 21 Qt 84 Kwb 138 Kmv 212 Kmg |
| 327 | 28-9479 | 20 Qt 76 Kwb 99 Kmv |
| 328 | 28-28512 | 18 Qt 138 Kmv+Kwb 195 Kmg |
| 329 | 28-20443 | 18 Qt 82 Kwb 124 kmv 231 Kmg |
| 330 | 28-8462 | 5 Qt 90 Kwb 110 Kmv 190 Kmg |
| 331 | 28-22529 | 9 Qt 58 Ket 128 Kwb 158 Kmv 187 Kmv+Kmg 257 Kmg |
| 332 | 28-16280 | 12 Qt 60 Ket |
| 333 | 28-15962 | 75 Ket |
| 334 | 28-26457 | 62 Ket 141 Kwb 186 Kmv 260 Kmg |
| 335 | 28-16335 | 70 Ket 150 Kwb 215 Kmv 265 Kmg |
| 336 | 28-26253 | 7 Qt 40 Ket 136 Kwb 188 Kmv 240 Kmg |
| 337 | 28-9803 | 63 Ket 65 Kwb |
| 338 | 28-31209 | 60 Ket 196 Kwb+Kmv 220 Ket |
| 339 | 28-4082, G | 40 Ket 120 Kwb 190 Kmv 280 Kmg 290 Kp |
| 340 | 28-21188 | 24 Tp 49 Ket 138 Kwb 184 Kmv 240 Kmg |
| 341 | 28-31351 | 29 Tp 57 Ket 137 Kwb 176 Kmv 247 Kmg |
| 342 | 28-23147 | 19 Tp 34 Ket 117 Kwb 176 Kmv 240 Kmg |
| 343 | 28-7997 | 9 Tp 35 Ket 55 Kwb |
| 344 | 28-31219 | 20 Tp 40 Ket 120 Kwb 158 Kmv 220 Kmg |
| 345 | 28-30475 | 16 Tp 49 Ket 123 Kwb 146 Kmv 252 Kmg |
| 346 | 28-21153 | 20 Tp 55 Ket 120 Kwb 160 Kmv 225 Kmg 280 Kmg +Kp |
| 347 | 28-7701 | 16 Tp 108 Kwb 122 Kmv 192 Kmg |
| 348 | 28-29752 | 14 Tp 27 Ket 112 Kwb 178 Kmv 240 Kmg |
| 349 | 28-21985 | 23 Tp 81 Kwb 140 Kmv 200 Kmg |
| 350 351 | 28-20780 28-8826 | 21 Tp 65 Kwb 103 kmv 110 kmg 19 Tp 70 Kwb 155 Kmv 185 Kmg |
| 352 | 28-27525 | 30 Tp 85 Kwb 100 Kmv 220 Kmg |
| 353 | 28-16500 | 28 Tp 99 Kwb 119 Kmv 180 Kmg |
| 354 | 28-17488 | 28 Tp 97 Kwb 121 Kmv 190 Kmg |
| 355 | 28-7592 | 32 Tp 80 Kwb 90 Kmv 144 Kmv+Kmg 199 Kmg |
| 356 | 28-4837 | 7 Tp 55 Kwb 69 Kmv 79 Kmg |
| 357 | 28-29913 | 75 Kwb 95 Kmv 160 Kmg |
| 358 | 28-30763 | 12 Tp 60 Kwb 100 Kmv 180 Kmg |
| 359 | 28-16751 | 35 Tp 90 Kwb 140 Kmv 180 Kmg |
| 360 | 28-31199 | 40 Tp 76 Kwb 134 Kmv 160 Kmg |
| 361 | 28-31200 | 39 Tp 76 Kwb 133 Kmv 158 Kmg |
| 362 | 28-29216 | 14 Tp 133 Kwb+Kmv 180 Kmg |
| 363 | 28-31201 | 40 Tp 76 Kwb 134 Kmv 160 Kmg |
| 364 | 28-11221 | 5 g 18 Tp 65 Kwb 89 Kmv 100 Kmg |
| 365 | 28-4846 | 10 Tp 64 Kwb+Kmv 70 Kmg |
| 366 | 28-23133 | 23 Kmv |
| 367 | 28-26010 | 12 Tp 30 Kmv |
| 368 | 28-853 | 7 Tp 13 Kmv 87 Kmv+Kmg 115 Kmg |
| 369 | 28-1828 | 22 f+Tp 60 Kmy 131 Kmg |
| 370 | 28-27595 | 8 Qt 25 Kmv 46 Kmg |
| 371 | SC 179 | 10 Qm 22 Kmg 53 Kp |
| 372 | DOT DR 84 | 13 w 21 Qm 58 Qal 60 Kp |
| 373 | DOT SB 269 | 25 Qm 62 Qal 86 Kp |
| 374 | DOT SB 307 | 26 Qm 57 Qal 81 Kp |
| 375 376 | DOT SB 346 DOT DR 92 | 37 Qal 60 Kp 37 w 53 Qal 60 Kp |
| 376 | DOT DR 92 DOT DR 97 | 12 w 51 Qal 60 Kp |
| 378 | DOT DR 100 | 7 w 10 Qm 30 Qal 60 Kp |
| 370 | DOI DK 100 | 1 v w 10 6m 20 6m 00 rb |

| Well Number | Identifier ¹ | Formations Penetrated ² |
|----------------|-------------------------|--|
| 379 | 28-30964 | 20 Qt 66 Kmg |
| 380 | 28-3660 | 41 Qal 45 Kmg 89 Kp |
| 381 | DOT DR 112 | 15 w 28 Qm 38 Qal 60 Kp |
| 382 | 28-456 | 30 Tp 110 Kmv+Kmg 155 Kmg 230 Kp |
| 383 | 28-7196 | 12 Tp 95 Kwb+Kmv 165 Kmg |
| 384 | 28-5083 | 130 Kmg 265 Kp, logged by F. J. Markewicz, 11/18/64 |
| 385 | DOT DR 130 | 7 w 11 Qm 60 Qal |
| 386 | 28-7100 | 42 Qt 212 Kp |
| 387 | 27-10794 | 17 Qwf 24 Kp |
| 388 | 28-28810 | 24 Qt 55 Kmg 120 Kp |
| 389 | DOT SB 243 | 16 Qm 40 Qal 130 Kp |
| 390 | 28-21-447 | 6 f 36 Qwf 39 sc |
| 391 | 28-29302 | 40 Qwf |
| 392 | 28-21-274 | 27 Qwf 59 wr? 61 r |
| 393 | 28-3645 | 22 Qal 70 Kmv 145 Kmg 299 Kp, logged by R. Mayer, 12/14/59 |
| 394 | 28-1601 | 13 Tp 92 Kmv+Kmg 232 Kmg 397 Kp |
| 395 | 28-16750 | 14 Qt 90 Kwb 150 Kmv 200 Kmg |

¹ Identifiers of the form 27-xxxx and 28-xxxx are N. J. Department of Environmental Protection well-permit numbers. Identifiers of the form 28-xx-xxx are N. J. Atlas Sheet grid locations of wells in the state well files or the N. J. Geological and Water Survey permanent note collection that do not have permit numbers. Identifiers prefixed by "DOT" are N. J. Department of Transportation test borings available at http://www.state.nj.us/transportation/refdata/geologic/. Identifiers prefixed by "SC" are test borings made in the 1930s for a proposed ship canal that are on file at the N. J. Geological and Water Survey. The identifier "core 1 (Southgate, 2010)" is a pollen core from: Southgate, E. W. B., 2010, Herbaceous plant communities on the Delaware River floodplain, New Jersey, during the mid-Holocene: Bulletin of the Torrey Botantical Society, v. 137, p. 252-262. A "G" following the identifier indicates that a gamma-ray log is available for the well.

²Number is depth (in feet below land surface) of the base of the unit indicated by an abbreviation following the number. Final number is the total depth of the well rather than the base of the unit. For example, "12 Tp 34 Kmg 62 Kp" indicates Tp from 0 to 12 feet below land surface, Kmg from 12 to 34 feet, and Kp from 38 to bottom of hole at 62 feet. Formation abbreviations and the corresponding drillers' descriptive terms (in parentheses) used to infer the formation are as follows:

Surficial Deposits: w=water (for test borings drilled in the Delaware River), f=fill, g=accretion gley overlying surficial unit (brown, gray silty clay, sandy clay), Qal=alluvium (gray, brown, black, yellow sand, gravel, silty sand), Qm=tidal-marsh deposits (black, brown, dark gray silt, sand, and clay with peat or organic matter), Qs=swamp and marsh deposits (black soil), Qt=terrace deposits, undivided (includes map units Qstl, Qstu, Qtl, Qtu, TQg) (yellow, gray, tan, brown, reddish-brown sand, clayey sand, gravel, silty sand, loam), Qwf=glaciofluvial deposit (includes map units Qwf and Qif) (brown, reddish-brown, yellowish-brown sand and gravel, sand), Tp=Pensauken Formation (yellow, yellowish-brown, orange-brown, red sand and gravel, sand, clayey sand, silty sand, loamy sand).

Cretaceous Formations: Ket=Englishtown Formation (yellow, white, orange, gray sand, with clay, wood, mica), Kwb=Woodbury Formation (gray, black clay, silty clay), Kmv=Merchantville Formation (green, gray, black clay, marl, silty clay), Kmg=Magothy Formation (gray, white fine sand, fine-to-medium sand, silt, clay, with wood, lignite, mica), Kp=Potomac Formation (white, yellow, brown, red, gray clay, sand, coarse sand, gravelly sand).

Pre-Cretaceous bedrock lithologies (not assigned to formations) are: cg=conglomerate, gb=gabbro, gn=gneiss, r=rock, q=quartzite, sc=schist, sh=shale, ss=sandstone. A "w" preceding the rock identifier indicates that is reported as "weathered" or "decomposed": wgn=weathered gneiss, wr=weathered rock, wsc=weathered schist, wss=weathered sandstone.

Units joined with a "+" cannot be separately identified in the driller's description. Units joined with an "&" are interbedded or intercalated. Queried units (for example, "Tp?") are inferred from inconclusive descriptions. Units are interpreted from drillers' or geologists' lithologic descriptions on records from the sources indicated in footnote 1. Well cuttings logged by N. J. Geological and Water Survey staff are noted with the geologist's name and date, if available. Units shown for wells may not match the map and sections due to variability in drillers' descriptions. For wells tapping surficial deposits, the depth of the screened interval (in feet below land surface) and yield (in gallons per minute) are shown following the log.