68 27-2475 0-6 subsoil (Qtt)

NEW JERSEY GEOLOGICAL SURVEY

INTRODUCTION Surficial materials in the Princeton quadrangle include wetland, hillslope, and windblown deposits, several generations of alluvial deposits, and weathered bedrock materials. These materials occur within a landscape sculpted by two major episodes of valley incision and dissection, accompanied by several major drainage shifts. The deposits are delineated on the map and sections and are described below. The large-scale erosional landforms are shown on figure 1, and small-scale erosional landforms such as fluvial scarps and rock-cut terraces are delineated on the map. The temporal relationships of the deposits and landforms are shown on the correlation chart.

DESCRIPTION OF MAP UNITS

not mapped.

af ARTIFICIAL FILL--Excavated bedrock and surficial material, and discarded man-made material (construction debris, cinders, ash, and trash). As much as 30 feet thick. Many small areas of fill, particularly in urban areas, are

Deposits in Modern Valleys (Holocene and Pleistocene)-This suite of deposits occurs in floodplains, terraces, wetlands, and on gentle valley-side slopes within modern valleys. It includes sediments of postglacial origin (units Qal, Qs, part of Qald and Qcal), of glacial origin (Qtt), and of periglacial origin (Qts, Qtc, Qaf, Qe, part of Qald and Qcal). The postglacial deposits are of Holocene and latest Pleistocene age. The glacial and periglacial deposits are primarily of Late Pleistocene age (the late Wisconsinan and, possibly, Illinoian

ALLUVIUM -- Silt, sand, clay, and pebble-to-cobble gravel. Moderately- to well-sorted, stratified. Contains variable but generally minor amounts of organic matter and peat. Color of fine sediment is brown and reddish brown in deposits derived from mudstone and sandstone, brown and yellowish brown in deposits derived from unit TQb and Coastal Plain formations. In deposits derived from mudstone and sandstone, the gravel consists of subangular to subrounded chips, flagstones, pebbles, and cobbles of mudstone and sandstone, with minor subrounded pebbles and cobbles of diabase and quartz, quartzite, and chert from units TQb and TQp in basins containing those units. In deposits derived from unit TQb and Coastal Plain formations, the gravel consists of subrounded to rounded quartz, quartzite, chert, and ironstone. Generally, overbank deposits of fine sediment overlie gravelly channel deposits. The overbank deposits are dominant in wide, gently-sloping floodplains; the channel deposits are dominant in narrow, steep floodplains. Total thickness of alluvium is as much as 20 feet in main valleys, 10 feet in tributary valleys (estimated). Unit may include some small deposits of colluvium (unit Qcsl) in narrow upland valleys

in the area north and west of Routes 206 and 27. SWAMP AND MARSH DEPOSITS--Black to brown peat overlying gray to dark gray silt and clay containing organic matter, Maximum thickness 10 feet (Neumann, 1976). Deposited in abandoned valleys after deposition of units Qtt and Qtc ceased and glacio-isostatic rebound began to reverse local drainages. These events likely occurred sometime between 14,000 and 10,000 years ago.

ALLUVIUM AND DIABASE BOULDER LAG--Silt, sand, minor clay and organic matter, dark brown to brown, moderately-sorted and weakly stratified; overlying and alternating with surface concentrations (lags) of rounded to subrounded diabase boulders and cobbles formed by winnowing of weathered diabase. Maximum thickness 10 feet (estimated). In places where valley gradients are steep,

unit consists of boulder fields with little fine sediment. COLLUVIUM AND ALLUVIUM, UNDIVIDED -- Gravelly, mudstone- or diabase-rich alluvium as in unit Oal, and colluvium as in units Qcsl and Qcd, filling the bottoms of narrow, steep tributary valleys in the upper Stony Brook basin and north of Mount Rose. Maximum thickness 10

Qe EOLIAN DEPOSITS--Reddish yellow to yellow silt and fine sand in the Millstone River, Stony Brook, and Shipetaukin Creek valleys, as much as 5 feet thick but generally less than 3 feet thick; and yellow fine sand, slightly micaceous, in Central Mercer County Park, as much as 10 feet thick. Well-sorted, nonstratified to weakly stratified. These windblown deposits were probably blown from the unvegetated surfaces of units Qtt and Qtc under periglacial conditions during the late Wisconsinan and, possibly, Illinoian, glaciations.

ALLUVIAL FAN DEPOSITS--Gravel, silt, and sand. Moderately-sorted, stratified. Color of fine sediment reddish brown to brown. Gravel consists of subangular to subrounded chips, flagstones, pebbles, and cobbles of mudstone and sandstone. As much as 10 feet thick

TRENTON GRAVEL--Sand, silt, pebble gravel, minor cobble

gravel. Well- to moderately-sorted, stratified. Color of fine sediment is reddish brown to reddish yellow, pale brown, and light gray. Sand is chiefly quartz and rock fragments, with some feldspar and a little mica. Gravel includes, in approximate order of abundance, gray and brown mudstone and sandstone, quartz, quartzite, chert, reddish brown mudstone and sandstone, and gneiss. The gray and brown mudstone and sandstone and the gneiss are derived chiefly from bedrock in the Delaware valley; the other clasts are derived chiefly from units TQb and TQp and from local bedrock. Some sandstone and gneiss clasts may have a thin weathering rind, otherwise the clasts are unweathered. Forms terraces 5 to 10 feet above the modern floodplain. As much as 30 feet thick.

This deposit is a northeasterly extension of the late Wisconsinan glaciofluvial gravel of the Delaware valley. It was termed the "Trenton gravel" by Cook (1880), and was included in the Cape May Formation by Bascom and others (1909a). The Cape May has recently been subdivided and redefined in its type area (Newell and others, 1989) and no longer includes glacial deposits, so the name is not used here. Owens and Minard (1975, 1979) renamed the deposits the "Spring Lake beds". They considered it to be an interglacial deposit of Sangamon age by correlation to an estuarine clay near Camden (the "Fish House clay" of Woolman, 1897). Mapping and field descriptions by Salisbury and Knapp (in Bascom and others, 1909b) show that the Trenton gravel in the Camden area forms a terrace cut into and below the Fish House clay, demonstrating that it is younger than Sangamon. Thus, the correlation and nomenclature of Owens and Minard (1979) is not used here, and the original nomenclature of Cook (1880) is

The Trenton gravel terrace extends northeasterly from the plain at Trenton, across the low Delaware-Millstone divide at Port Mercer, and down the lower Millstone valley to Manville, about 16 miles northeast of Princeton, where it merges with glaciofluvial deposits in the Raritan valley. From Trenton to Manville the surface of the terrace maintains a nearly constant elevation of about 60 feet. When adjusted for the effects of postglacial crustal rebound, which has raised the northern end of the terrace relative to the southern end, an even slope down the Delaware valley, across the divide at Port Mercer, and down the Millstone valley, results (Stanford, 1990). This reconstruction indicates that during the late Wisconsinan glaciation (which reached its maximum about 20,000 years ago), with the earth's crust depressed to the north by the weight of the glacier, glacial meltwater in the Delaware valley was partially diverted down the lower Millstone and deposited unit Qtt. A similar diversion may have occurred during the Illinoian glaciation (about 150,000 years ago), but Illinoian sediments within unit Qtt have not been

observed in the Princeton quadrangle. Upon deglaciation, crustal rebound gradually flattened the terrace gradient northeast of Trenton. Eventually, drainage across the divide at Port Mercer stopped, and valleys V2, V3, and V4 (fig. 1) were abandoned. A radiocarbon date of 10,430±1160 years (I-16,554, plotted on map) on basal peat directly overlying the Trenton gravel on the divide at Port Mercer is a minimum date for valley abandonment.

TERRACE DEPOSITS, COASTAL PLAIN PHASE-Sand and pebble gravel, minor silt and clay. Well-sorted, stratified. Color of fine sediment is pale brown, yellow, light gray, gray, and brown. Sand is chiefly quartz but includes some mica and glauconite. Gravel is chiefly quartz and quartzite, with minor chert and ironstone and a little gray mudstone near contact with unit Qtt. The sand is derived from Coastal Plain formations and unit TOb; the gravel is derived chiefly from unit TQb. Forms terraces 5 to 10 feet above the modern flooplain. As much as 30 feet thick. Qtc terraces are on grade with Qtt terraces, and Qtc interfingers with Qtt along their contact. These relationships indicate that the units are largely contemporaneous, and that Qtc was deposited under periglacial conditions as deposition of Qtt in the Assunpink-Millstone lowland caused valley aggradation in tributary valleys.

TERRACE DEPOSITS, MUDSTONE AND SANDSTONE PHASE--Gravel, sand, and silt. Moderately-sorted, stratified. Color of fine sediment is reddish brown to brown. Sand is chiefly quartz and rock fragments, with some feldspar and a little mica. Gravel consists of subangular to subrounded chips, flagstones, pebbles, and cobbles of chiefly mudstone and sandstone, with some quartz, quartzite, and chert in the deposit along Shipetaukin Creek and minor diabase in the Stony Brook deposits. Forms terraces 5 to 10 feet above the modern floodplain. As much as 15 feet thick (estimated). Age relationships and origin same as unit Qtc.

Deposits that Predate the Formation of Modern Valleys (Late Miocene?-Early Pleistocene?)--These sediments were deposited by former streams not directly related to the modern drainage network. They are extensively dissected and deeply weathered. They are generally preserved as erosional remnants on hilltops and divides, although they are still continuous in the southeastern corner of the map, where they are thick.

Bascom and others (1909a) included these deposits with the Pensauken Formation, which is the lower yellow gravel of the Delaware valley south of Trenton (Salisbury and Knapp, 1917). Owens and Minard (1975, 1979) divided these deposits into an upper and a lower unit. They correlated the lower unit with the Pensauken Formation, and the upper unit with the Bridgeton formation, which is the upper vellow gravel of the Delaware valley south of Trenton (Salisbury and Knapp, 1917), based on projection of the base elevation of the units from their type areas in southern New Jersey. Because the yellow gravel at East Trenton Heights forms a distinctly lower surface than the yellow gravel elsewhere in the quadrangle, the nomenclature and correlations of Owens and Minard (1975, 1979) are adopted here.

The ages of the Bridgeton and Pensauken are uncertain. Berry and Hawkins (1935) describe plant fossils from the Bridgeton near New Brunswick (about 15 miles northeast of Princeton) that they consider to be of Pleistocene age. Owens and Minard (1979) assign both units a Late Miocene age based on correlation to units in the Delmarva Peninsula, although this correlation cannot be physically demonstrated because the deposits are not continuous across Delaware Bay. Pollen from a black clay bed within the Bridgeton, exposed in a gravel pit near Plainsboro (about 4 miles east of Princeton), includes cool-temperate to coldtemperate species and a few exotic pre-Pleistocene species (G. Brenner, written communication), an assemblage suggesting a Pliocene age. These data indicate that it is possible that the deposits span the period from the Late Miocene to the

PENSAUKEN FORMATION--Sand and pebble gravel, minor cobble gravel and silt. Well-sorted, stratified. Color of fine sediment is yellow, reddish yellow, and light gray. Sand is chiefly quartz with some weathered feldspar. Gravel is chiefly quartz and quartzite with some chert, ironstone, and decomposed sandstone and mudstone. As much as 65 feet thick. Forms a terrace with a surface altitude of 55-60 feet that is incised about 40 feet below the Bridgeton surface. May have been deposited by Assunpink Creek and other local tributaries that dissected and reworked the Bridgeton deposit after diversion of the

BRIDGETON FORMATION .- Sand and pebble gravel, minor cobble gravel, silt, and clay. Well-sorted, stratified. Color of fine sediment is yellow, reddish yellow, and light gray. Sand is chiefly quartz with some weathered feldspar Gravel is chiefly quartz and quartzite with some chert, ironstone, and decomposed sandstone and mudstone. As much as 90 feet thick. Deposited by a large river flowing southwesterly along the inner edge of the Coastal Plain from the New York City area to the Delmarva Peninsul (Salisbury and Knapp, 1917; Owens and Minard, 1979). Deposition ceased and the Bridgeton valley was abandoned when the Bridgeton river was diverted seaward in the New York City area, possibly during Late Pliocene or Early Pleistocene glaciation (Stanford, 1990).

Correlation Chart

Hillslope Deposits (Pliocene?-Late Pleistocene)--These sediments form aprons at the base of hillslopes. They are deposited by downslope movement of weathered bedrock material, chiefly under periglacial conditions when permafrost may be present, vegetation is scant, and slopes are unstable. The deposits are predominantly of Late Pleistocene age, but may be older in the

subsurface parts of unit Qcd in a few places protected from Pleistocene erosion. MUDSTONE AND SANDSTONE COLLUVIUM, LOWER PHASE--Poorly-sorted, nonstratified sediment composed of angular to subangular chips and flagstones of mudstone and sandstone in a matrix of reddish brown sandy claves silt. Tabular clasts generally have a strong slope-paralle orientation. Forms aprons with surfaces that grade distally to heights of 5 to 10 feet above the modern floodplain, or that grade distally to the surface of unit Qts. As much as 15 feet thick (estimated). This deposit grades to and

PHASE--Sediment as in unit Qcsl. Forms aprons on rockcut benches in the upper Stony Brook valley with bases 10 to 30 feet above the modern floodplain. As much as 15 feet thick (estimated). These deposits predate final incision of Stony Brook and deposition of units Qcsl and Qts. DIABASE COLLUVIUM--Poorly-sorted, nonstratified sedi-

cobbles of diabase in a matrix of reddish yellow clayey silt to slightly sandy silt. Maximum thickness 20 feet Qcdl DIABASE COLLUVIAL LAG--Surface concentrations of with small amounts of organic matter. Maximum thickness

Weathered Bedrock Material (pre-Cretaceous?-Holocene)--Sediment formed by mechanical and chemical weathering of bedrock. These materials are predominantly of Quaternary age but may be of Tertiary age in places where they have been protected from Pleistocene erosion--for example, on parts of upland erosion surfaces \$1 and \$2 (fig. 1) and subjacent to unit TQb. Subjacent to Cretaceous deposits (unit K), the weathered rock material may be partly of

> WEATHERED MUDSTONE AND SANDSTONE--Poorlybedrock); few to many angular to subangular chips, flagstones, pebbles, and cobbles of sandstone (some sand (on sandstone bedrock). Sand is mostly quartz but bedrock, which crops out primarily south and east of Small outcrops are common in streambanks, stream beds

WEATHERED DIABASE, BOULDERY PHASE--Poorlysorted, nonstratified material consisting of some to many subangular to subrounded diabase cobbles and boulders in a matrix of yellow, reddish yellow, brown, and light gray clayey silt to slightly sandy silt. Generally less than 10 feet thick. Occurs on sloping surfaces on Rocky Hill. Unit Odwt is thin, bouldery broken-rock rubble with little

matrix material, on steep slopes. WEATHERED DIABASE, SILTY PHASE--Moderatelysorted, nonstratified material consisting of few subangular to subrounded, partially decomposed pebbles and cobbles of diabase in a matrix of reddish yellow, red, and brown clayey silt. Maximum thickness 15 feet (estimated). Occurs on flat to very gently sloping surfaces on the

WEATHERED SCHIST AND GNEISS--Poorly-sorted, nonstratifed to weakly stratified material consisting of some subangular to subrounded pebbles and cobbles of partially decomposed gneiss or schist in a matrix of ellow, brown, yellowish brown, white, olive yellow, olive brown, and light gray micaceous sand to clayey sand (on gneiss bedrock) and micaceous clayey silt to silty clay (on

summit of Rocky Hill.

schist bedrock). As much as 70 feet thick. CRETACEOUS DEPOSITS, UNDIFFERENTIATED-Sand, silt, and clay of the Raritan and Magothy Formations of Cretaceous age. Variably oxidized to white, yellow, pink, red, brown, pale brown, and gray. As much as 200 feet thick in the southeastern corner of the

where gradational or feathering, dotted where concealed by

interfingers with unit Qts and so is, in part, contemporaneous with units Qts, Qtt, and Qtc. MUDSTONE AND SANDSTONE COLLUVIUM, UPPER

ment composed of subrounded to rounded boulders and

diabase by spring scepage and surface runoff on hillslopes.

rounded diabase boulders and thin, discontinuous, reddish yellow to brownish yellow clayey silt to slightly sandy silt, 10 feet (estimated). Occurs on gently-sloping aprons formed by winnowing of fine material from weathered

Cretaceous or older age (sections AA', BB', CC'). At the surface the weathered bedrock units include discontinuous colluvium generally less than 3 feet thick.

> sorted, nonstratified to weakly stratified material consisting of some to many angular to subangular chips of red to gray mudstone in reddish brown, red, reddish yellow, and yellow silty clay to clayey silt (on mudstone decomposed) in yellow, pale brown, white, light gray, reddish brown, red, and reddish yellow clayey sand to silty includes some mica, weathered feldspar, and rock fragments. Generally less than 10 feet thick on mudstone bedrock, which crops out primarily north and west of routes 206 and 27; as much as 80 feet thick on sandstone routes 206 and 27. On sandstone bedrock weathering occurs preferentially in certain beds (chiefly in coarsegrained arkosic sandstones), and zones of unweathered ock may overlie or interbed with zones of weathered rock. and man-made excavations on mudstone bedrock but are less common on sandstone bedrock.

quadrangle (Zapecza, 1989).

Fluvial scarp of postglacial age-Line at top of scarp, ticks on Fluvial scarp of Early to Middle Pleistocene age--Line at top

MAP SYMBOLS

Contact--Dashed where approximately located, short-dashed

of scarp, ticks on slope.

Fluvial scarp of Bridgeton age (Late Miocene-Early Pleistocene)--Line at top of scarp, ticks on slope. Marks northwest border of unit TQb. Artificial excavation scarp--Line at top of scarp, ticks on slope.

 Crest of low bedrock ridge--Line at crest of low (generally less than 6 feet high) bedrock-controlled ridge, parallel to strike of beds. Drawn from aerial photographs. ★ Inactive sandstone quarry

Marks extent of sandstone quarries.

Shallow topographic basin--Line marks rim of basin, dot marks bottom of basin. Height of closure generally less than 10 feet. Of probable periglacial origin (Wolfe, 1953). Rock-cut terrace--Flat surfaces cut on shale bedrock by Stony Brook, about 10-15 feet above the modern floodplain. Cut in

the Early to Middle Pleistocene, before final incision of Stony Brook to form the modern floodplain. Bodies of water--Not shown on base map. Qe/TOb Thick eolian sand overlying Bridgeton Formation--Eolian sand continuous and generally greater than 6 feet thick.

(Qe)/Qsw Thin eolian loam overlying weathered mudstone and sandstone--Eolian loam generally continuous but less than 3

• 47 Well or boring with log in table 1--Location judged to be accurate to within 100 feet 047 Well or boring with log in table 1--Location judged to be accurate to within 500 feet.

from well and boring data (table 1).

Elevation of base of Bridgeton and Pensauken formations--

Contour interval 20 feet. Shown only in southeast part of

map where the base of the units is not exposed. Mapped

Site of radiocarbon-dated peat--With laboratory number and age. Sample taken at a depth of 5 feet, directly on top of unit

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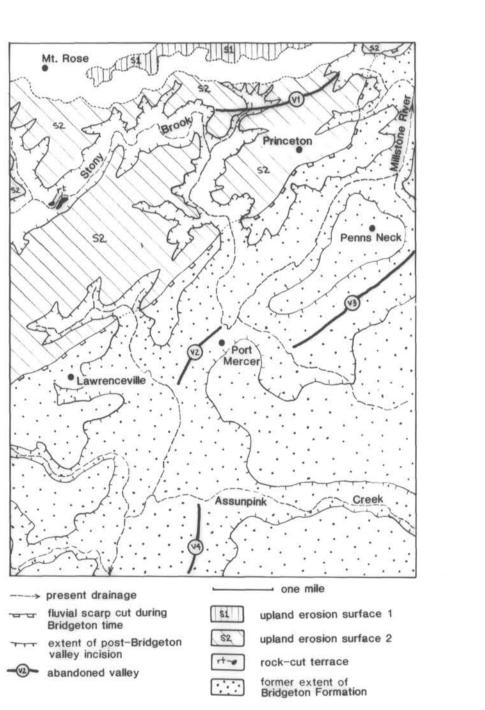
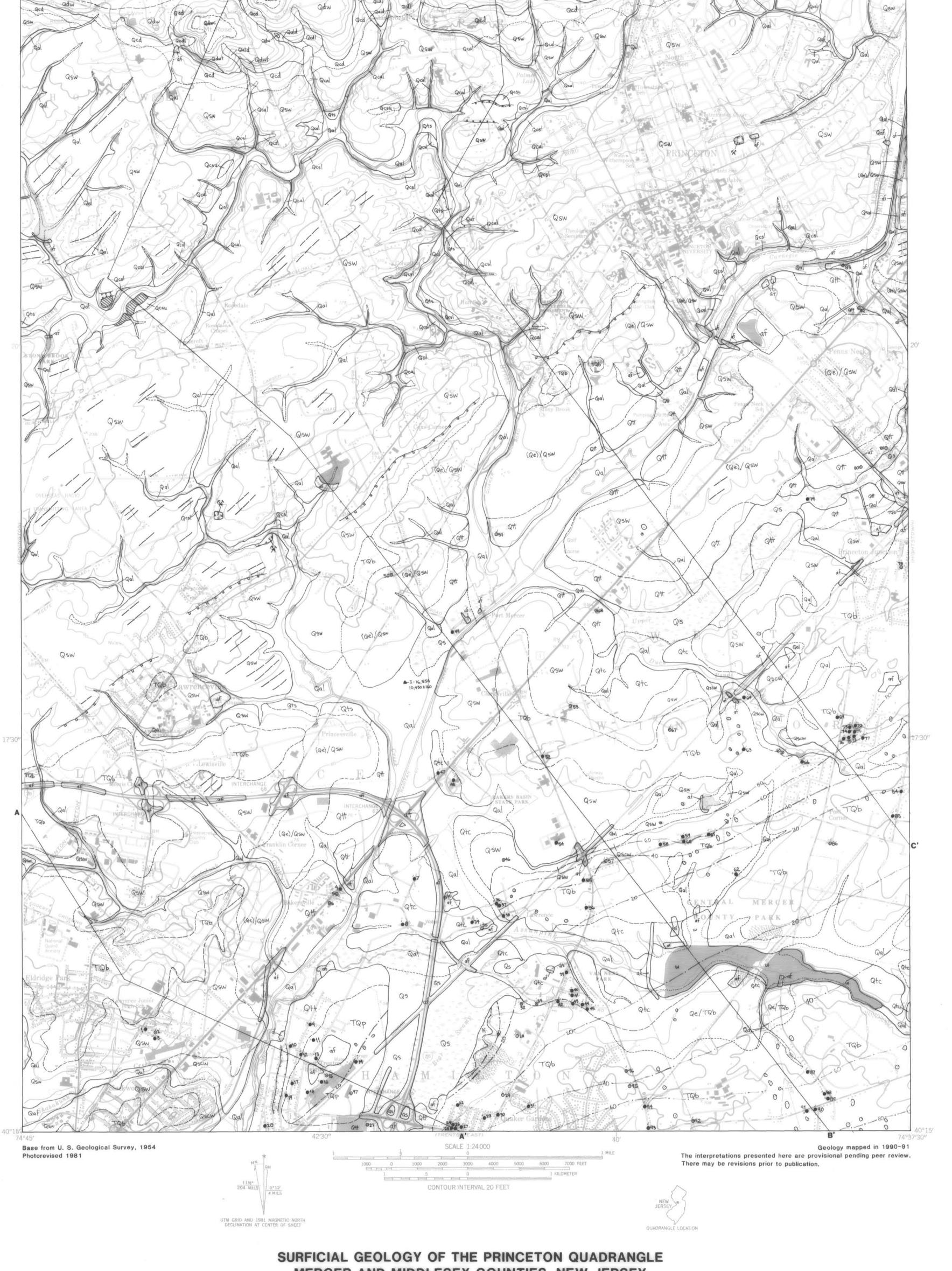
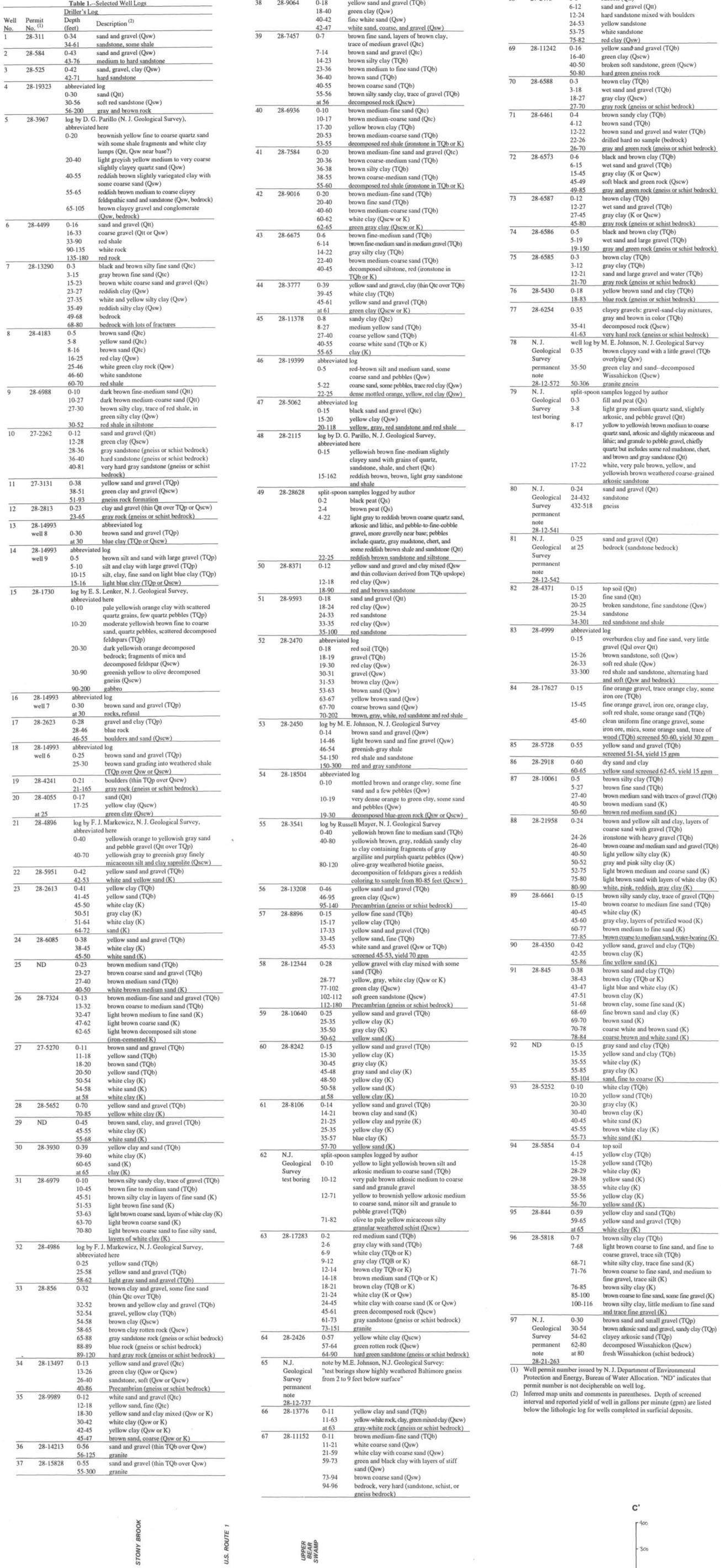


Figure 1. -- Erosional landforms of the Princeton quadrangle. Upland erosion surfaces are flat to very gently sloping surfaces on summit and interfluve areas that truncate bedding and layering in the underlying bedrock. Abandoned valleys are stream-cut valleys that do not carry through-flowing streams today. The extent of post-Bridgeton valley incision is drawn approximately, at the location of slope inflections from undissected upland areas on erosion surface 2 and remnants of unit TQb to steeper slopes on the sides of incised valleys.



MERCER AND MIDDLESEX COUNTIES, NEW JERSEY

Scott D. Stanford 1993

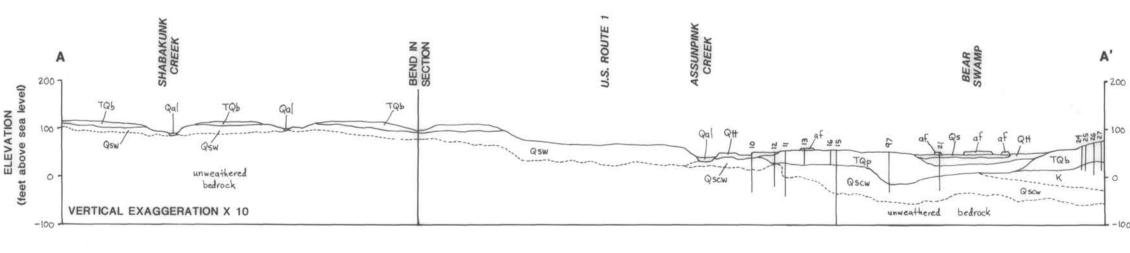


28-9064 0-18 yellow sand and gravel (TQb)

10 to 30 ft of incision in main valleys; V4 possibly cut by Assunpink Creek and rt cut diversion of Bridgeton river; southerly-draining Assunpink and northerly-draining Millstone established D U and S2 cut; V1 cut by east-flowing proto-Stony Brook 200 to 300 ft of walley incision; widespread dissection \$1 cut K

*(S1, S2, rt, V1, V2, V3, V4 are shown on fig. 1; fluvial scarp symbols are identified unde "Map Symbols" and are shown on the map.)

Erosional Landforms and Events*



VERTICAL EXAGGERATION X 1

unweathered bedrock

unweathered bedrock

unweathered bedrock

unweathered bedrock VERTICAL EXAGGERATION X 10