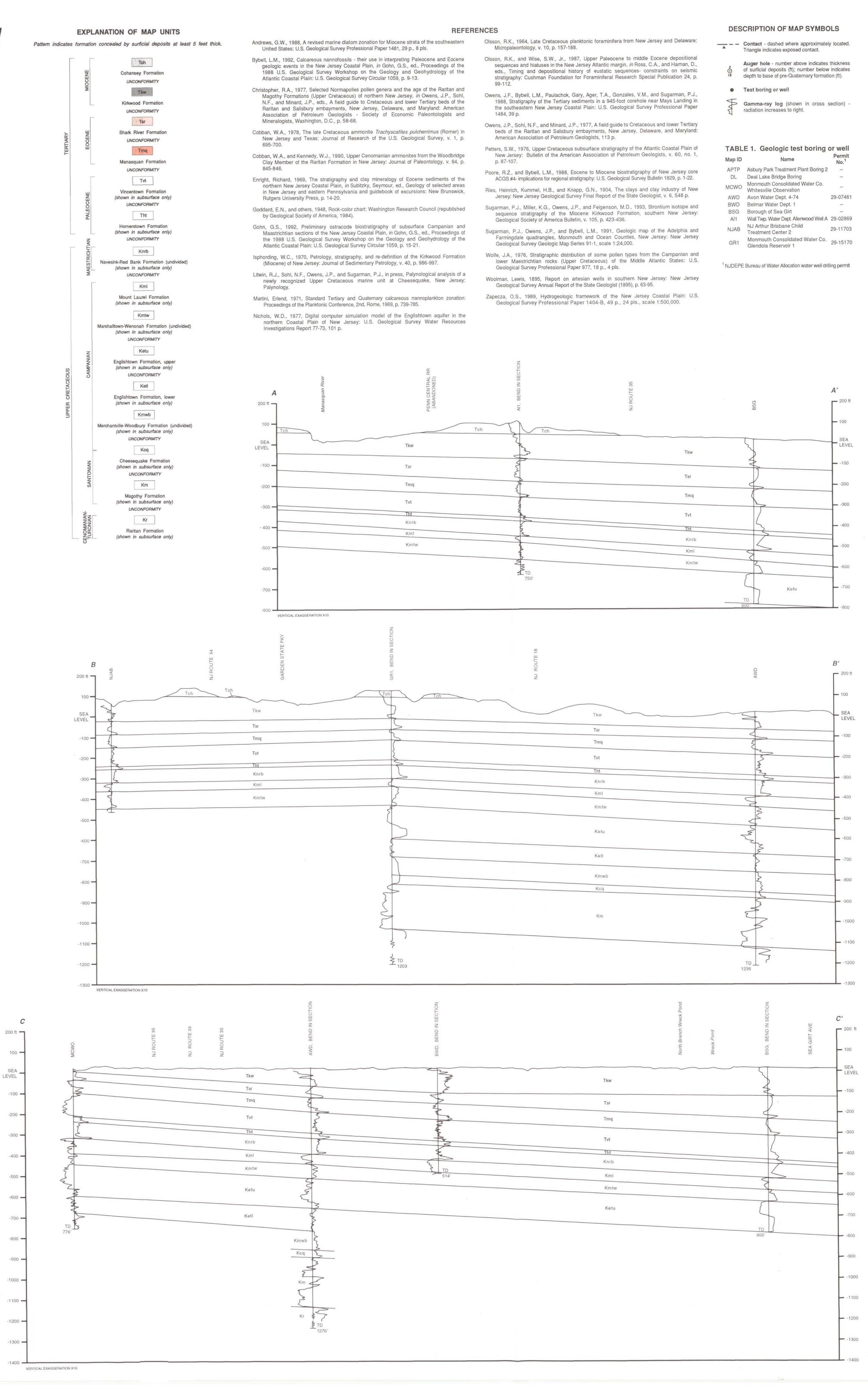
LITHOLOGY THICKNESS (FT) "Surficial" refers here to all post-Cohansey deposits. These are predominantly alluvial, colluvial, and paralic sands and gravels. They are shown on older maps as belonging to the Bridgeton, Pensauken, and Cape May Formations, but the use of these names for deposits distant from the type localities and of uncertain relationship to the type deposits is questionable. SURFICIAL **DEPOSITS** Alluvial deposits in present-day stream channels and adjacent lowlands are not shown on this map. Older alluvium forms terraces adjacent to stream channels and caps some hills. The terrace alluvium consists predominantly of quartz sand and gravel. Adjacent to the Manasquan River it is more than 20 feet (6.1 m) thick. The alluvium capping hills consists of orange and brown gravelly sand up to 10 feet (3 m) thick. The clasts are predominantly quartzite, with a maximum length of 3 inches (7.6 cm). Colluvium derived from mass wasting of the Cohansey Formation is widespread at lower elevations in river valleys and the coastal lowlands. It covers extensive areas of the eastern half of the quadrangle. Paralic deposits include Holocene beach sands and estuarine sediments. Some of the lakes adjacent to the coast previously were estuaries. In the Deal Lake Bridge boring, 55 feet (16.8 m) of Holocene material consists of fluvial, estuarine, and nearshore marine facies. Radiocarbon dating of peat and woody material yielded a date of 10,190 years for the base of this \Holocene transgressive sequence, and 3,080 years for the middle of the section (Meyer Rubin, U. S. Geological Survey, written communication, 1990). Quartz sand, yellowish-gray to pinkish-gray and grayish-orange to dark yellowish-orange, medium- to very coarse-grained, with occasional pebbles. Cross-stratified (trough and planar-tabular). Dominantly an orthoguartzitic sand with traces of weathered feldspar and chert. Detrital heavy minerals may be abundant (2 to 3 percent) and ilmenite dominates among the opaque minerals; to a lesser extent zircon and sillimanite dominate the nonopaque minerals. Kaolinite is the dominant clay-sized mineral in the sand matrix and locally replaces feldspar clasts at the base of paleochannels. Ophiomorpha burrows as **FORMATION** much as 1.5 inches (3.8 cm) in diameter are common in certain beds. The map area is northeast of the main outcrop belt. The formation is extensively eroded, limited to elongate outliers only a few feet or tens of feet thick. Exposures are poor except in excavations because of the loose, sandy nature of the formation. Numerous pits offered excellent exposures of the Cohansey Formation during the mapping of the quadrangle, but many have subsequently closed. The basal contact of the formation is placed at the unconformity between cross-bedded, medium-to-coarse sand of the Cohansey and massive, fine-grained micaceous sand of the Kirkwood Formation. No dateable material has been recovered from the Cohansey in this quadrangle. Owens and others (1988) consider the Cohansey as middle Miocene, owing to the similarity of its palynomorphs to those of the Kirkwood Formation. Recent strontium isotope age estimates for the upper part of the Kirkwood Formation (Sugarman and others, 1993) indicate that the Cohansey Formation is no older than approximately 12 m.y. (middle Miocene). Sand, typically light-colored, interbedded with and overlying dark-gray or brown clay-silt. The sand is quartzose with small amounts of feldspar and mica (mostly muscovite). It is typically massive, orange, yellow, or gray, micaceous, burrowed, and has extensive iron oxide (Leisegang) banding. The detrital heavy minerals are dominated by the opaques, especially ilmenite, with lesser amounts of nonopaques including zircon, staurolite, garnet, tourmaline, and rutile. The lowermost clay-silt, termed the Asbury Clay (Ries and others, 1904) or Asbury Park Member of the Kirkwood Formation (Isphording, 1970) is a dark, peaty, massive to laminated clay-silt with occasional interbeds of fine sand. Pyrite is common in the dark beds. Finely dispersed clay minerals FORMATION include kaolinite, illite, and illite/smectite. This member is exposed in the Shark River and Manasquan River valleys. A reworked, coarse glauconite-quartz sand with granules at the base of the Kirkwood is typically a few feet thick and rests unconformably upon the Shark River Formation. Woolman (1895) reported the diatom Actinoptychus heliopelta from the Kirkwood Formation in a well at Asbury Park. This places the Kirkwood within the lower Miocene East Coast Diatom Zone 1 of Andrews (1988). Sugarman and others (1993) obtained a strontium isotope age estimate of 19.2 to 22.6 Ma ±0.5 m.y. for this zone. Clay-silt to very fine quartz sand, calcareous (where unweathered), grayish-olive-green and light-to-dark greenish gray, massive to thick-bedded and extensively burrowed. Botryoidal glauconite, fine to coarse, as much as 20 percent in some intervals, is disseminated in a dominantly clay-silt matrix. Clay minerals include illite, illite/smectite, kaolinite, and minor clinoptilolite. Layers 1 to 2 feet (0.3 to 0.6 m) thick are cemented with iron oxide and contain casts and molds of Venicardia antiquata. Exposures of the Shark River are sparse and very weathered. Thin, 1 to 2 foot (0.3 to 0.6 m) ledges and float blocks occur in the Shark River upstream from Schoolhouse Road; thicker exposures form banks along the upstream branches of Deal Lake. The contact SHARK RIVER with the underlying Manasquan Formation is unconformable (Sugarman and others, 1991; Owens and others, 1988; Poore and Bybell, 1988). Because of **FORMATION** the limited exposure and deep weathering of Shark River and Manasquan outcrops, the contact is inferred. Its location near Deal Lake is based on updip projections from contacts identified in split spoon samples from Deal Lake and the Asbury Park Wastewater Treatment Plant. The Shark River is middle Eocene (Bybell, 1992; Poore and Bybell, 1988). In the N. J. Geological Survey Allaire State Park corehole on the adjacent Farmingdale quadrangle, the unit contained calcareous nannoplankton zones NP 14 to 16 (Sugarman and others, 1991). Split spoon samples from Shark River Park and the Asbury Park Wastewater Treatment Plant contained the same zones. Clay-silt, dusky-yellow-green to pale-olive and grayish-green, extensively burrowed, calcareous (where unweathered), massive to thick-bedded, grading upward into very fine quartz sand. Clay minerals include illite, illite/smectite, and minor clinoptilolite. Fine glauconite sand is commonly dispersed weather out of the formation and are washed ashore. The Manasquan grades downward in some places into a quartz-glauconite sand. Enright (1969) and Olsson and Wise (1987) subdivide the Manasquan and Shark River into members. Subdivision is impracticable in this quadrangle because of the limited exposures and discontinuity of the members. The Manasquan is lower Eocene (Owens and others, 1988; Poore and Bybell, 1988; Sugarman and others, 1991). Split spoon samples from the Deal Lake Bridge borehole included calcareous nannoplankton zones NP 9 (upper part) to NP 13. Sand, grayish-olive to dusky-yellow-green to greenish, dark-greenish, and olive-gray. Mostly quartz with minor amounts of feldspar. Glauconite composes 2 to 3 percent of the sand in the upper part; it increases to as much as 30 percent toward the base. Silt and fine sand similarly increases towards the base. Pyrite nodules, phosphate pebbles, and mica occur locally. Clay-sized minerals in the matrix include illite, illite/smectite, kaolinite, and calcite. Fossils are numerous and include bryozoan, mollusk, and echinoid fragments and microfossils. The quartz sand is an updip facies which changes sharply downdip to a grayish-olive-green, massive, micaceous, burrowed clay-silt with thin beds of very fine quartz and glauconite sand. The sand facies is thickest in the outcrop where it reaches more than 140 feet (43 m) south of Long Branch (Zapecza, 1989). Based on its foraminifera, the Vincentown in this area is upper Paleocene (Olsson and Wise, 1987). Glauconite sand, clayey, massive-bedded, dusky-yellowish-green to dusky-green; greenish-black where unweathered. Glauconite grains are mainly medium to coarse in size and botryoidal. Contains 1 to 2 percent fine- to very coarse-grained quartz sand, phosphate fragments, pyrite, and lignite. Based on its calcareous nannofossils, the Hornerstown is lower Paleocene (Danian) (Bybell, 1992). Sand, glauconite, slightly quartzose, clayey, greenish-black. Unconformably underlies the Hornerstown Formation and overlies the Mount Laurel NAVESINK-RED BANK In outcrop the Navesink Formation and the Red Bank Formation form an unconformity-bounded, coarsening-upward sedimentary sequence consisting of FORMATION (undivided) a basal glauconite sand (Navesink Formation), a middle silt, and an upper quartz sand (Red Bank Formation). Downdip but to the west of the Asbury Park quadrangle, the sand pinches out and the silt changes facies to a glauconite sand which is lithologically indistinguishable from the underlying Navesink Formation. The nannofossils Nephrolithus frequens and Lithraphidites quadratus indicate an upper Maestrichtian age. Sand, quartz, fine- to coarse-grained, glauconite (2 to 5 percent), extensively burrowed, slightly micaceous and feldspathic, commonly interbedded with MOUNT LAUREL thin layers of dark clay-silt. Olive-gray to dark-greenish gray where unweathered. The transition from the Mount Laurel to the underlying Marshalltown-Wenonah Formations (undivided) is gradational and is marked by an increase in mica, a decrease in grain size, and the appearance of more massive clay-silt beds in the Marshalltown-Wenonah. Variation in the thickness of the formation is common, Based on macrofossils, the Mount Laurel is assignable to the upper Campanian Exogyra cancellata zone (Owens and others, 1977). Based on foraminifera, it is assigned to the lower Maestrichtian Rugotruncana subcircumnodifer zone (Petters, 1976). MARSHALLTOWN-WENONAH Glauconite sand, greenish-black, massive-bedded, extensively burrowed, with fine-grained quartz sand and silt (Marshalltown), grading upward into a FORMATION (undivided) thick, very silty, micaceous sand (Wenonah). The Marshalltown-Wenonah is recognized in the subsurface by a small to large gamma spike at the base of the Marshalltown passing into a relatively flat, high-intensity pattern above. Undivided due to the thinness of the Marshalltown Formation (approximately 10 ft or 3 m) and its lithologic similarity to the Wenonah Formation. The Marshalltown-Wenonah Formations (undivided) have been assigned to the upper Campanian, based on the occurrence of the ammonite Trachyscaphites pulcherrimus in the outcropping Wenonah Formation in the Marlboro quadrangle (Cobban, 1978) and the occurrence of the Globotruncana calcarata zone in the Marshalltown Formation in the Woodstown quadrangle (Olsson, 1964). Clay-silt to very fine quartz sand, glauconitic, dark-greenish-gray, micaceous and lignitic. Grades upward into a fine-grained sand interbedded with thin, dark-gray, micaceous, woody, clay-silt. The sand is chiefly quartz; less than 10 percent consists of feldspar, rock fragments, and glauconite. Thin-walled FORMATION, UPPER mollusks are common in the lower clay-silt facies. Defined on the gamma-ray log by a thick, high-intensity clayey unit at its base and a thinner, low-intensity sand at its top. Gohn (1992) tentatively assigns a middle Campanian age to this formation (termed S3 cycle in Gohn, 1992) based on its ostracode content. Quartz sand, feldspathic, micaceous and lignitic, fine- to medium-grained, medium- to dark-gray, interstratified with dark-greenish-gray, carbonaceous clay-silt. Glauconite, mica, and lignite are common constituents. The top of the unit is identified on the gamma ray log as below the basal gamma spike of the upper part of the formation; the base is defined in this quadrangle as the base of the first sand above the Merchantville-Woodbury Formations (undivided). In FORMATION, LOWER southeastern Monmouth and northeastern Ocean counties, the Englishtown Formation had been subdivided into an upper and lower sand facies divided by a clay-silt facies (Nichols, 1977; Zapecza, 1989). The clay-silt facies and upper sand facies are shown here as the upper Englishtown Formation. Clay-silt with very fine sand, dark-gray, finely micaceous, with occasional lenses of finely disseminated pyrite, lignite, and siderite. Bedding is massive to finely laminated with alternating layers of very fine sand and clay-silt, occasionally cross-bedded. Grades downward into an intercalated, thick-bedded sequence of glauconitic sand and silt and micaceous clayey silt. The glauconitic sand is grayish-olive, greenish-black, or dark greenish-gray; the clay-silt MERCHANTVILLE-WOODBURY is shades of black and gray. Quartz and glauconite are the major sand components; feldspar, mica (colorless and green), and pyrite are minor constituents. Siderite-cemented layers are common. The formation is highly bioturbated and contains zones of broken calcareous mollusks. Wolfe (1976) assigned the Woodbury to the lower Campanian based on palynomorph assemblages; the Merchantville is also lower, but not lowermost, Campanian based on the ammonite Scaphites hippocrepis III (Owens and others, 1977). Clay-silt, glauconite (20 percent maximum), brownish to dark gray, massive, burrowed, and very micaceous. Grades to light-gray, silty, very fine to fine sand at top, generally laminated where not extensively burrowed. Lower contact with the Magothy Formation and upper contact with the Merchantville Formation are irregular, extensively burrowed, and contain layers of pale-brown to grayish-brown siderite concretions as much as 1 foot (0.3 m) in diameter. The Cheesequake is a recently named formation consisting of beds formerly included in the Merchantville Formation. It contains an uppermost Santonian to lowermost Campanian pollen assemblage in the outcrop and subsurface (Litwin and others, in press). It crops out in the South Amboy quadrangle. Intercalated quartz sand and clay, thin- to thick-bedded. Sand is light- to medium-gray or brownish-gray; clay is olive-black to grayish-black. Bedding is horizontal (laminated) and cross-stratified. The sand is fine to very coarse, well sorted within each bed, predominantly quartz, and includes minor feldspar and mica. Pyrite-cemented and pyrite-coated sand concretions are common. Carbonaceous material is abundant in beds as much as 0.5 feet (0.15 m) FORMATION The Magothy is Upper Cretaceous (Santonian) based on Zone V pollen at the Freehold core hole (Christopher, 1977). Clay-silt, massive. The upper part of the formation is olive-gray to dark greenish-gray, the middle part light- to medium-gray, and the lower part reddish-brown, light- to dark-gray, reddish-purple, red, and brown. The upper part contains interbeds of fine to coarse sand, occasionally cross-bedded, with broken shell material. Mica, lignite, pyrite, and siderite are common accessories. Siderite forms layers 0.25 to 0.50 inch (0.6 to 1.3 cm) thick. The Raritan is assigned to the upper Cenomanian (Upper Cretaceous) pollen Zone IV - the Complexiopollis-Atlantopollis zone at the Freehold core hole and the Marlboro MUA well. The ammonite Metoicoceras bergquisti establishes an upper Cenomanian age at outcrops in Sayreville (Cobban and Kennedy, **FORMATION** Column not to scale



GEOLOGIC MAP OF THE ASBURY PARK QUADRANGLE,

Reviewed by O.S. Zapecza

Cartography by Z. Allen-Lafayette



MONMOUTH AND OCEAN COUNTIES, NEW JERSEY

CONTOUR INTERVAL 20 FEET DEPTH CURVES AND SOUNDINGS IN FEET-DATUM IS MEAN LOW WATER

DECLINATION AT CENTER OF SHEET