

## INTRODUCTION

The Wanauke quadrangle is located in north-central New Jersey and includes parts of Morris, Passaic, and Bergen Counties. Surficial deposits in the quadrangle include glacial deposits of late Wisconsinan age and postglacial deposits. The glacial deposits include till, which occurs in drumlins and as a discontinuous veneer draping bedrock uplands; and stratified sand, gravel, silt, and clay partially filling valleys. The stratified glacial deposits were deposited in glacial lakes and outwash sheets and are as much as 180 feet thick. The till may be as much as 200 feet thick in drumlins. Postglacial deposits overlie the glacial deposits in places and include swamp and marsh sediment, alluvium, and alluvial fan deposits. They are generally less than 25 feet thick.

## PREGLACIAL DRAINAGE

Preglacial drainage in the Wanauke quadrangle generally followed the routes of the present drainage. However, two major dislocations resulting from the deposition of stratified sediment in valleys are identified. In the Wanauke Valley south of the Wanauke axis of the preglacial valley, which has been deepened by glacial erosion, underlies the valley now occupied by Post Brook while the present Wanauke flows in the more narrow valley to the east now occupied by Lake Inez. The second dislocation is in the Pompton Lakes, where the preglacial Ramapo Valley, also deepened by glacial erosion, continues southwestward from Pompton Lake to join the Wanauke, while the present Ramapo River flows southerly through a gap in Packanack Mountain in the Pompton Plains quadrangle.

## GLACIAL ADVANCE

The orientation of drumlins and striations on uplands indicates that late Wisconsinan ice flowed to the south and southeast as it advanced across the quadrangle. This flow direction is also indicated by the location of plucked bedrock outcrops, which generally occur on south- or southeast-facing slopes; and polished and aligned bedrock outcrops, which generally occur on north- and northwest-facing slopes. Additionally, till (Qt) is generally thicker and more continuous on northwest-facing slopes, suggesting lodging of till during advance of ice to the south and southeast. One exception to this ice-flow direction is in the southeastern corner of the quadrangle, east of the Ramapo Mountains. Here, striations and landforms in areas adjacent to the quadrangle indicate that advancing ice was moving to the southwest. Ice in this part of the quadrangle was in the fringe of a major south-flowing lobe in the Newark Basin lowlands, which lie east and south of the Ramapo Valley.

The distribution of pebbles in till corroborates these ice-flow directions. Lithologies of pebbles in till were counted at 20 sites in the quadrangle (table 2). Till north and west of the Ramapo Valley contains a mix of gneiss pebbles derived from local bedrock, and gray shale and sandstone, some purple quartzite and conglomerate, and a few dolomite and chert pebbles. The shale, sandstone, quartzite, conglomerate, dolomite, and chert were transported southward and easterly from the Green Pond outlier, a belt of Paleozoic rock to the northwest of the quadrangle, and to a lesser extent, from the Wallkill Valley farther to the northwest. Scattered erratic boulders of dolomite, indicated by a special symbol on the map where observed, also indicate southerly to southeasterly transport. Anomalous high percentages of gray shale and sandstone at two places (sites 23 and 26) are in till that may have been deposited by melting of upper parts of the glacier containing far-traveled clasts. At site 23 this till overlies till with a 96-percent gneiss-pebble content. The predominance of local clasts in this till indicates that it was deposited at the base of the glacier. Till east and south of the Ramapo Valley contains significant percentages of red sandstone and conglomerate, and some basalt, from the Newark Basin. These rock types are absent in till north and west of the Ramapo Valley, indicating that ice did not flow westward from the Newark Basin lowlands.

## RECESSIONAL DEPOSITS

Late Wisconsinan ice advanced to the Terminal Moraine, approximately 12 miles southeast of Butler. Retreat of ice from the Terminal Moraine began approximately 20,000 years ago (Cotter and others, 1986) and was accompanied by a reorientation of ice flow to the southwest, parallel to the regional topographic grain. This ice-flow direction is recorded by the northwest-southeast orientation of recession ice margins reconstructed from the positions and elevations of stratified deposits. The positions of these recession ice margins and the extent of their associated recession deposits are shown on figure 1. The reorientation may have been the result of thinning of ice and consequent enhanced topographic control of ice flow.

As the ice front receded northward across the quadrangle, stratified sand, gravel, and silt were deposited in glacial lakes and fluvial plains. Glacial lakes formed where the retreating ice front blocked north-draining valleys or where open drainage down valleys was blocked by ice or by previously-deposited glacial sediment. Stratified sediment deposited in glacial lakes includes detritic and lacustrine-fan sand and gravel deposited at the receding ice margin; and lake-bottom fine sand, silt, and minor clay which settled onto the lake floor away from the ice margin. These units are included in either the "Upland Lacustrine and Ice-Contact Deposits" or the "Fluvial and Fluvial-Lacustrine Deposits" groups in the description of map units.

Fluvial plains were deposited from successive recession ice-margin positions in valleys where drainage was not blocked or where lakes were either filled with sediment, drained when sediment dams were eroded, or drained when ice dams melted or collapsed. Fluvial deposits include pebbly sand to boulder gravel and in places overlie detritic and lake-bottom sand at depth. These units are grouped as "Fluvial and Fluvial-Lacustrine Deposits" in the description of map units.

Small, short-lived glacial lakes dammed by the retreating ice in north-draining valleys on uplands also contain detritic sand and gravel and, in places, thick, sandy, bouldery diamict that may be either flow deposits from the glacial surface or collapsed lacustrine deposits. These units are grouped as "Upland Lacustrine and Ice-Contact Deposits" in the description of map units. Also included in this group are several diamict deposits (Qkm, Qhm1, Qhm2) occupying small bowl-shaped valleys bounded on their north sides by high hills. These deposits may have formed when downwasting of the glacier caused ice blocks to become isolated and stagnant in the lee of steep bedrock hills. These stagnant ice blocks melted and released sediment to form the observed diamict deposits. An exception to this rule is the diamict deposit just east of Skyline Lakes (Qsk). Here, diamict occurs in several arcuate ridges that rise well above the level of the small north-draining valley they occupy. This morphology suggests that the deposit was emplaced by active ice.

Lithologies of pebbles in recession deposits were counted at 56 sites (table 2). These counts show that the recession deposits have pebble mixes that are similar to the adjacent till or that are slightly more enriched in local bedrock than the till. This pattern suggests that the recession sediment was derived from erosion of till and bedrock by meltwater flowing beneath or along the edge of the glacier. An exception to this pattern are the deposits in the Ramapo Valley. Even though these deposits rest on sandstone and conglomerate of the Newark Basin, they contain between 61 and 83 percent gneiss and only 1 to 2 percent Newark Basin rocks. This distribution indicates that these deposits were supplied primarily by meltwater descending the east slope of the Ramapo Mountains rather than by subglacial meltwater from the valley bottom.

## DEGLACIATION HISTORY

The first set of recession ice margins in the quadrangle is marked by thick lacustrine and diamict deposits in the north-draining tributary valleys on the south side of the Pequannock Valley (Qml, Qmd1, Qmd2, Qmd3, Qmd4). Similar deposits on the north side of the valley (Qhm, Qhm1, Qhm2, Qhm3, Qhm4) appear the main valley was occupied by stagnant ice. These ice margins parallel the deep, narrow east-west trending valley of the Pequannock River and may have formed when melting back of the ice front was slowed by the much thicker ice occupying this deep valley.

As ice retreated from this valley, open drainage down the Pequannock was established and outwash was deposited along the Pequannock (Qpw) and two south-sloping tributary valleys (Qow, Qob). Farther east the large fluvial unit (Qpp) forming the Pompton Plains south of the Wanauke quadrangle was deposited from an ice margin near the quadrangle boundary, and a small lacustrine unit (Qil) was deposited in a ponded valley north of Pines Lake. Somewhat later, the first fluvial-lacustrine unit in the Wanauke Valley (Qw1) was deposited. Also at this time, outwash was being deposited in a south-sloping tributary valley north of Pompton Lakes (Qpl) and lacustrine deposits (Qpl) were accumulating in a glacial lake dammed by ice to the north of Pecknack Mountain, just east of Oakland.

When the receding ice crossed the divide between the Pequannock and the Wanauke at the head of the valley north of Bloomingdale, a proglacial pond formed in the north-sloping valley south of Twilliger Lake. Unit Qm2 was deposited in this pond. Soon thereafter, as continued recession opened drainage eastward into the Wanauke Valley, outwash deposits Qw and, somewhat later, Qpp formed. Farther west, Qgl was deposited in a proglacial pond south of Gordon Lakes.

At approximately the same time in the Ramapo Valley, fluvial-lacustrine unit Qrm1 was deposited, probably from an ice margin just outside the quadrangle to the northeast. Later, Qrm1 was incised by a later unit (Qrm2) deposited from an ice margin farther up-valley in the Ramsey quadrangle.

The next major recession ice margin is marked by thick lacustrine deposits in north-sloping tributaries on the south side of West Brook Valley (the earlier deposits Qmp, Qms, Qm1, followed by Qw1, Qw2, Qw3, Qw4). By the head of fluvial-lacustrine unit Qw2 in the Wanauke Valley, by the ice contact deposit at Skyline Lakes (Qsk), and by two small lacustrine units in north-sloping tributaries to the Wanauke (Qw5, Qw6). Like the Pequannock Valley, the West Brook Valley is a deep, east-west trending trough. The thick deposits in this valley, like those in the Pequannock, may represent slowed retreat of the ice margin in response to the thick section of ice occupying the valley.

Retreat of the ice margin from the West Brook position opened drainage down West Brook, allowing deposition of outwash (Qw6). As the remainder of the quadrangle was deglaciated, lacustrine units (Qm1, Qm2, Qm3, Qm4, Qm5, Qm6, Qm7) were deposited in several north-sloping valleys. Fluvial units were deposited along the south-sloping valleys of Cusaw Brook (Qc1, Qc2), Ringwood Creek (Qw1, Qw2), and the Wanauke River (Qw3, Qw4). Where (Qw3, Qw4). Where steeply-sloping meltwater channels debouched into valleys, coarse gravel was deposited in fans and channels (Qm1).

The far northwest corner of the quadrangle was occupied by a portion of a glacial lake partially flooding the north-draining valley of Belchers Creek, which is just northwest of the edge of the map. This lake drained across the Wanauke-Pequannock divide in the Newfoundland quadrangle, approximately 2 miles west of Midvale Lake, at an elevation of approximately 850 feet (Stanford, 1991). Deltas (Qw3, Qw4) were deposited from two ice margins in this lake.

## POSTGLACIAL DEPOSITS

After deglaciation, deposition of silt and clay, followed by peat and muck, began in ponds, marshes, and swamps (Qs). Most of these wetlands occupy shallow basins glacially eroded in bedrock on uplands. Several wetlands in the quadrangle occupy floodplains and former lake-bottoms in the valleys.

When meltwater deposition ended, postglacial alluvium (Qal) began to accumulate along low-gradient reaches of streams and rivers. These deposits are generally thinner, finer-grained, and of smaller extent than the meltwater deposits and in some places contain organic matter. Two small alluvial fans (Qf) of cobble- to-boulder gravel on the north side of the Pequannock valley near Asplawa are exceptions to the generally fine-grained character of postglacial deposits. These fans were probably deposited shortly after deglaciation by erosion of the loose sediment in deposits Qhm and Qhm.

## REFERENCES

Cotter, J. F. P., Ridge, J. C., Evenson, E. B., Sevon, W. D., Sirk, Les, and Stuckenrath, Robert, 1986, The Wisconsinan history of the Great Valley, Pennsylvania and New Jersey, and the age of the Terminal Moraine in Cadwell D.H., (ed.), The Wisconsinan Stage of the first geological district, eastern New York, N. Y. State Museum Bulletin 455, p. 22-49.

Stanford, S. D., 1991, Surficial geologic map of the Newfoundland quadrangle, Passaic, Morris, and Sussex Counties, New Jersey: N. J. Geological Survey Geologic Map Series 91-3, scale 1:24,000.

Wakman, S. A., Schulhoff, H., Hickman, C. A., Cordon, T. C., and Stevens, S. C., 1943, The peats of New Jersey and their utilization: N. J. Department of Conservation and Development Geologic Series Bulletin 55, Part B, 278 p.

TABLE 1-- Selected well logs

Well No.	Permit No.	Driller's Log		Interpretation	
		Depth (feet below land surface)	Description	Depth (feet below land surface)	Unit
1	in file at N.J. Geological Survey	9-9	medium to fine sand, gravel	9-9	Qd1
		10-42	fine sand, gravel	10-70	Qd1
		40-76	fine sand, gravel	10-70	Qd1
		70-118	medium to fine sand and gravel	10-70	Qm1b
2	in file at N.J. Geological Survey	20-20	coarse sand, gravel, large stones	0-23	Qm1
		22-26	fine sand, gravel		
		25-28	gray clay, fine sand, gravel	23-19	Qm1b
		189-19	gray clay and sand, gray clay gravel	119-23 to 119-22	Qd1 (includes boulder)
3	20-5756 and 20-5117	0-24	coarse sand and gravel	0-24	Qm1
		100-171	fine sand, gravel, boulders	100-171	Qm1b
4	20-4584	171-228	fine sand, gravel, boulders	171-228	Qm1b
		10-38	sand and gravel, gray clay	0-16	Qm1
5	20-4579	10-38	gray clay	10-20-13	Qm1b
		10-38	sand, silt, light gray clay	0-16	Qm1
		45-55	heavy gray clay, fine gravel	10-20-13	Qm1b
		50-56	fine gravel	0-16	Qm1
6	20-5134	0-18	fine sand, gravel, boulders	0-18	Qm1
		18-70	gray clay, fine sand, gravel	0-60	Qm1b (includes boulder)
		70-84	fine sand, gravel, fine clay	0-60	Qm1b
		80-90	coarse sand, gravel, fine clay	0-60	Qm1b
7	20-11802	95-119	fine sand, gravel	95-110	Qm1b
		0-13	fine sand, gravel, boulder		
		10-115	fine sand, gravel, boulder		
		114-175	fine sand, clay, gravel	110-134	Qm1b (includes boulder, debris, debris, and silt; boulder is remainder of unit of Qm1)
8	20-5236	100-300	fine sand and gravel	200-300	Qm1b
		45-55	fine sand, gravel, boulders	0-43	Remainder of unit of Qm1b
		100-144	fine sand, gravel, boulders	100-144	Qm1b
		100-144	fine sand, gravel, boulders	100-144	Qm1b
9	20-6979	0-50	coarse sand, gravel, boulders	0-50	Qm1b (includes remainder of unit of Qm1)
		100-144	fine sand, gravel, boulders	100-144	Qm1b
10	20-13066	84-105	fine sand and gravel	84-105	Qm1b (includes remainder of unit of Qm1)
		100-100	gray clay	0-100	Qm1b (includes remainder of unit of Qm1)
			Unit: boulders	100-200	Qm1b

1. New Jersey Department of Environmental Protection and Energy water well drilling permit

TABLE 2-- Lithology of pebbles in surficial units

Site	Unit	Number of pebbles	Lithology (percentages of pebbles)				Remarks
			Greenish <sup>(1)</sup>	Gray Quartzite and Sandstone <sup>(2)</sup>	Purple Quartzite and Conglomerate <sup>(3)</sup>	Red Sandstone <sup>(4)</sup>	
1	Qs	102	10	50	26	0	(6)
2	Qs	121	8	53	9	0	(6)
3	Qs	140	19	44	16	0	(6)
4	Qs	141	22	54	23	0	(7)
5	Qs	205	48	48	17	0	(6)
6	Qs	159	63	21	16	0	(6)
7	Qs	151	74	27	26	0	(6)
8	Qs	154	65	24	11	0	(6)
9	Qs	154	76	22	2	0	(6)
10	Qs	154	74	25	0	0	(6)
11	Qs	154	74	25	1	0	(7)
12	Qs	213	42	52	6	0	(6)
13	Qs	243	18	40	40	0	(6)
14	Qs	184	25	54	21	0	(6)
15	Qs	125	0	16	84	0	(6)
16	Qs	139	72	22	6	0	(6)
17	Qs	154	66	37	27	0	(6)
18	Qs	154	48	39	13	0	(6)
19	Qs	189	36	22	32	0	(6)
20	Qs	189	36	22	32	0	(6)
21	Qs	189	60	37	3	0	(6)
22	Qs	121	34	34	0	0	(6)
23	Qs	147	22	48	6	0	(6)
24	Qs	142	44	48	8	0	(6)
25	Qs	142	70	22	6	0	(6)
26	Qs	84	5	63	32	0	(6)
27	Qs	117	28	28	44	0	(6)
28	Qs	118	45	11	44	0	(6)
29	Qs	165	61	36	3	0	(6)
30	Qs	162	62	36	2	0	(6)
31	Qs	143	63	36	1	0	(6)
32	Qs	143	49	39	0	0	(6)
33	Qs	126	78	19	3	0	(6)
34	Qs	126	78	17	4	0	(6)
35	Qs	134	75	20	5	0	(6)
36	Qs	134	75	20	5	0	(6)
37	Qs	129	77	19	4	0	(6)
38	Qs	129	77	23	4	0	(6)
39	Qs	141	71	28	1	0	(6)
40	Qs	136	73	17	0	0	(6)
41	Qs	136	73	17	0	0	(6)
42	Qs	147	68	29	3	0	(6)
43	Qs	153	42	42	1	0	(6)
44	Qs	151	80	14	3	0	(6)
45	Qs	129	67	17	1	0	(6)
46	Qs	127	83	13	2	2	(6)
47	Qs	127	87	17	14	0	(12)
48	Qs	136	74	18	7	0.5	(12)
49	Qs	154	64	34	0.5	0	(12)
50	Qs	164	55	37	8	0	(12)
51	Qs	161	57	39	12	0	(6)
52	Qs	161	57	39	12	0	(6)
53	Qs	144	52	37	10	0	(6)
54	Qs	151	21	11	0	0	(6)
55	Qs	159	65	18	17	0	(6)
56	Qs	157	68	28	4	0	(6)
57	Qs	155	85	23	20	0	(6)
58	Qs	111	77	19	4	0	(6)
59	Qs	137	68	26	8	0	(6)
60	Qs	129	76	11	13	0	(6)
61	Qs	139	65	4	30	0	(7)
62	Qs	129	78	16	6	0	(6)
63	Qs	148	68	26	6	0	(6)
64	Qs	153	65	26	9	0	(6)
65	Qs	142	77	17	3	0	(6)
66	Qs	142	84	24	0	0	(6)
67	Qs	157	76	22	2	0	(6)
68	Qs	147	68	24	8	0	(6)
69	Qs	205	71	19	10	0	(6)
70	Qs	97	30	15	0	56	(14)
71	Qs	88	68	14	3	0	(6)
72	Qs	131	77	17	5	1	(6)
73	Qs	158	66	26	10	1	(16)
74	Qs	157	64	26	10	0	(6)
75	Qs	155	75	23	1	0	(6)
76	Qs	152	75	23	1	0	(6)

- (1) Includes all igneous and metamorphic rock of Precambrian age.  
(2) Includes Belted Sandstone, Monksville Shale, and minor amounts of Marlburg Formation from northwestern of the quadrangle.  
(3) Includes chiefly Gneiss, Monksville Shale, and Green Pond Conglomerate.  
(4) Includes sandstone, siltstone, shale, and conglomerate of the Passaic Formation.  
(5) Includes 1 percent chert and 1 percent dolomite.  
(6) Includes 0.5 percent chert and 0.5 percent dolomite.  
(7) Includes 1 percent dolomite.  
(8) Count is in a pre-advance fluvial gravel underlying till (exposed in excavation). Not a map unit.  
(9) Includes 1 percent chert.  
(10) Upper unit of till.  
(11) Includes 0.5 percent basalt.  
(12) Includes 0.5 percent chert.  
(13) Includes 0.5 percent basalt.  
(14) Includes 1 percent basalt.

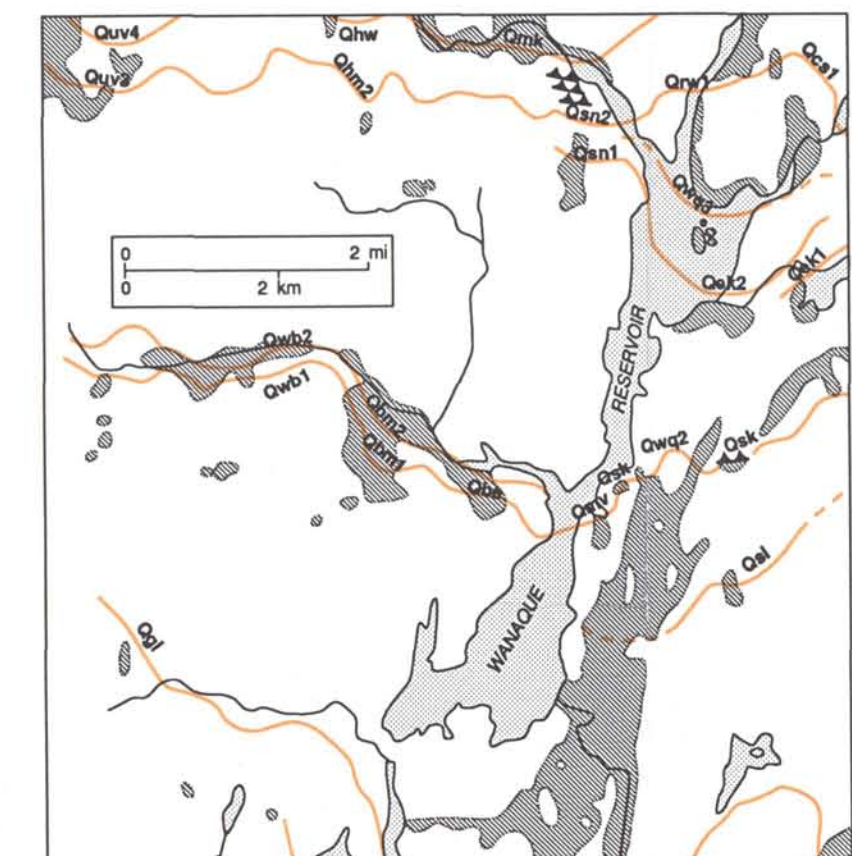


Figure 1. Recessional ice margins (in color) and associated deposits (in black) in the Wanauke quadrangle. Barbed lines indicate ice-margin ridges.

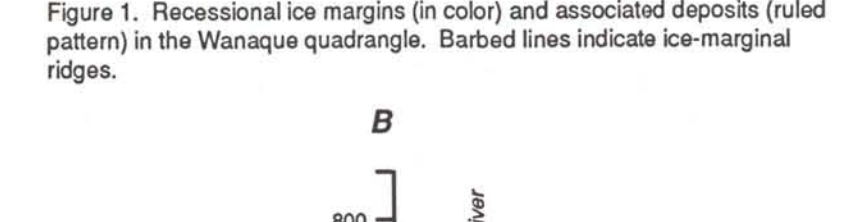


Figure 2. Vertical exaggregation X5. The figure shows a cross-section of the Wanauke quadrangle, with labels for various units and locations. The vertical axis represents elevation in feet above mean sea level, ranging from 0 to 800.

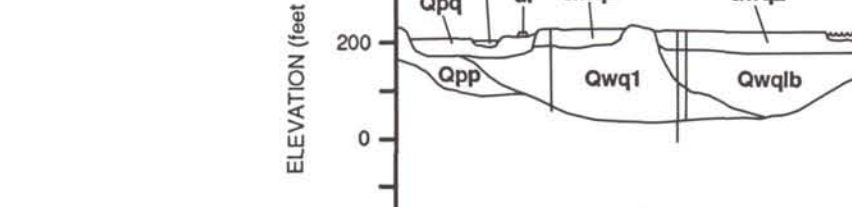


Figure 3. Vertical exaggregation X5. The figure shows a cross-section of the Wanauke quadrangle, with labels for various units and locations. The vertical axis represents elevation in feet above mean sea level, ranging from 0 to 800.

